

Results from a Worldwide Proficiency Test on the Determination of Carbonic Species Concentration in Natural Waters

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ABSTRACT

The results of an international inter-laboratory proficiency test for the determination of carbonic species concentration are presented. Eight laboratories performed the analysis of twelve water samples (four synthetic waters, one lake water, four geothermal waters, one seawater, and two petroleum waters) by two methods: (a) individual laboratory analytical procedure and (b) acid-base titration curves in tabular form. From the titration curves the concentrations of carbonic species are calculated using the Hydrologist method, Geochemist's method and initial pH and total alkalinity method. The titration curves of all laboratories yield very similar results (i.e., concentration of carbonic species for each sample). In contrast, the individual laboratory reported values showed high dispersion about the mean. This implies that there are some laboratories, which have problem in their concentration calculation procedure. To apply the Hydrologist and/or Geochemist methods it is necessary to locate always two equivalence points (e.g., NaHCO₃EP and H₂CO₃EP) even for samples which have pH lower than that of NaHCO₃EP. In addition, the backward titration curve with NaOH after complete removal of CO₂ is strictly necessary to decide for the applicability of the Hydrologist method or the Geochemist method. In cases where the complete analyses of species that contribute to the alkalinity are known, the initial pH and total alkalinity method is appropriate.

1. INTRODUCTION

The geochemistry of natural systems is much more complex than that of any synthetic system studied in the laboratory. When we extend our geochemistry knowledge acquired through laboratory experiments to natural systems, we face certain limitations. Mankind adopts an empirical approach based on creating an analytical database for similar systems around the world. This implicitly demands for reliable analytical measurements in every laboratory. Extensive efforts are in progress worldwide to create reference materials (standards) for each chemical species and to calibrate analytical techniques with such materials to achieve accuracy and consistency in the analytical database on natural geological systems.

Carbon dioxide (CO₂) plays a fundamental role in governing the geological and environmental processes on Earth. The distribution of carbonic species (H₂CO₃, HCO₃⁻ and CO₃²⁻) in natural waters permits the examination of CO₂ exchange between atmosphere and water bodies, the evaluation of buffering mechanisms and the definition of their acid-base neutralizing capacity (Stumm and Morgan 1996). Similarly, measurements of total alkalinity and carbonic alkalinity are of great importance in analyses of ocean, marine, lake and river waters, and other samples. The knowledge of aquatic CO₂ chemistry is fundamental for the removal of anthropogenic CO₂ from atmosphere (NASA 2012), incrustation of calcite in the reservoir and production wells during geothermal exploitation. Recently, Torres-Alvarado et al. (2012) presented the propagation of analytical uncertainty (error) in the determination of carbonic species concentration of separated water at the weir box in the pH calculation of geothermal reservoir fluids.

Verma (2013) performed the statistical analysis of the results of inter-laboratory comparisons of chemical analysis of geothermal waters, conducted under the quality assurance and quality control programs of International Association of Geochemistry and Cosmochemistry (IAGC) and International Atomic Energy Agency (IAEA). There was an overall uncertainty of $\pm 13\%$ and difficulty in defining the appreciable improvement in the analytical quality in the successive inter-laboratory comparisons. This is probably due to the existence of systematic errors in the measurements from some laboratories. Specifically, there were some problems with sampling and analytical procedures for SiO₂ and HCO₃⁻. In case of SiO₂, Verma et al. (2012, 2015) found that one of the major reasons for high dispersion in the measured values among laboratories was associated with minor differences in the individual analytical procedures of the participating laboratories. van Geldern et al. (2013) presented an inter-laboratory

comparison of the determination of isotopic composition ($\delta^{13}\text{C}$) of the total dissolved inorganic and organic carbon of different types of waters.

To further investigate the analytical challenges with carbonic species analyses the present study was initiated. It consists of a worldwide inter-laboratory comparison for the determination of carbonic species concentration in 12 water samples (four synthetic waters, one lake water, four geothermal waters, one seawater, and two petroleum waters), distributed among eight participating laboratories.

2. PREVIOUS INTER-LABORATORY COMPARISON FOR CARBONIC SPECIES

Ellis (1976) reported the results of first inter-laboratory comparison of chemical analysis of seven water samples, including a geothermal water sample. However, he reported only the statistics of the dataset without the individual measured value of each chemical parameter. The IAEA conducted a series of inter-laboratory comparisons for the chemistry of geothermal waters within the framework of the project, "Coordinated Research Program on the Application of Isotope and Geochemical Techniques in Geothermal Exploration" (Giggenbach et al. 1992, Gerardo-Abaya et al. 1998, Alvis-Isidro et al. 1999, 2000, 2002, Urbino and Pang 2004) and concluded that the dispersion (scatter) in the measured values among laboratories was a consequence of water sample alteration during transportation and storage.

Figure 1 presents a summary of the statistics of HCO_3^- concentration determinations, analyzed in thirteen water samples during the IAEA inter-laboratory comparison program (Verma, 2014). There is wide spread in the concentration values, reported by the participating laboratories (Figure 1a). Figure 1b presents a whisker plot between HCO_3^- concentration and RSD (relative standard deviation) for each sample after removal of outliers (Verma 2004). There is a considerably high analytical uncertainty for the measurements of HCO_3^- ; however, each participating laboratory had good reproducibility. Figure 1c shows the time evolution of RSD. It looks that there is improvement in the analytical quality for HCO_3^- with time. However, if we also include in the analysis the concentration of carbonic species in the samples (i.e., a relation between HCO_3^- , RSD and time), the improvement in the data quality is not well defined. Verma and coworkers (Verma 2004, 2005, 2008, Verma et al. 2010, 2014) initiated a project to understand theoretical and experimental aspects of the analysis of carbonic species in natural waters. The volumetric acid-base titration is the only reliable method for the determination of carbonic species concentration in natural waters. From an analytical point of view there exist four principle types of natural water: 1.) rainwater containing low concentration of dissolved species and low alkalinity, 2.) ground and surface water which contains mainly carbonic alkalinity, 3.) geothermal water which has carbonic and other alkalinities, and 4.) petroleum waters which have high concentration of salts and no contact with the present atmospheric CO_2 .

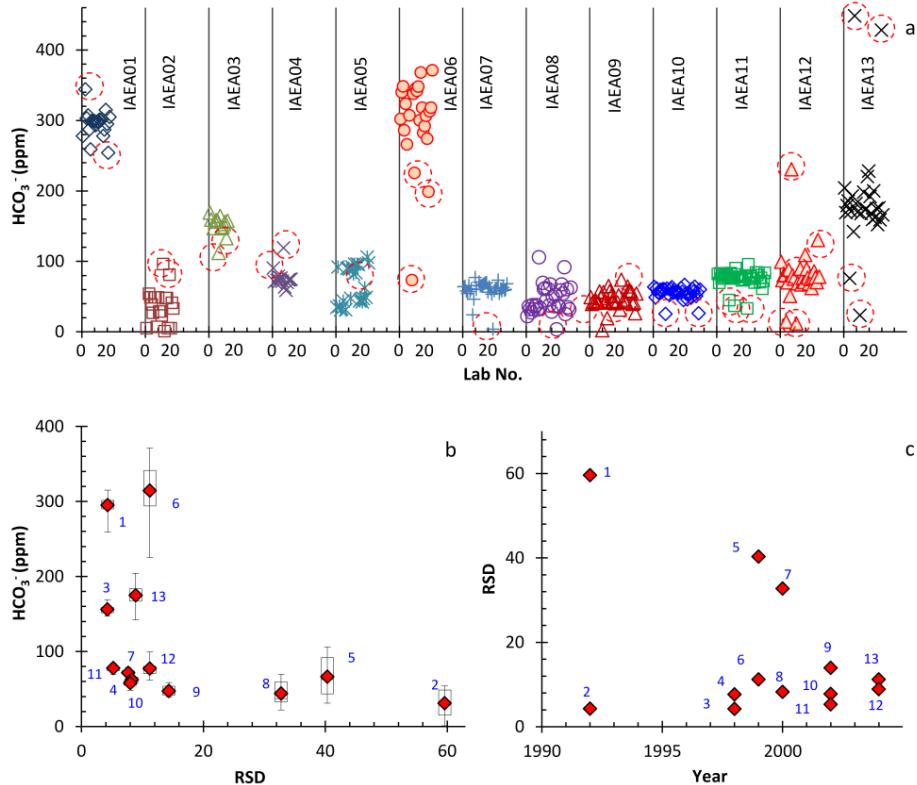


Figure 1: IAEA Inter-laboratory comparison of HCO_3^- in geothermal waters (after Verma 2014): (a) Behavior of individual laboratory data, (b) Whisker plot for the concentration and RSD (% relative standard deviation) and (c) Time evolution of RSD. The outliers according to the statistical procedure (Verma 2004) are marked with dashed circle.

3. SAMPLE PREPARATION AND CARBONIC SPECIES DETERMINATION METHODS

Verma et al. (2014) described the characteristics of 12 water samples (coded as IIE26 to IIE37), distributed during the present inter-laboratory comparison. The first four samples (IIE26 to IIE29) were prepared in the laboratory at Instituto de Investigaciones Electricas (IIE) by dissolving certain amounts of Na_2CO_3 , NaCl , KCl and NaOH in deionized water. Additionally, there were one

natural lake water (IIE30), four geothermal waters (IIE31 to IIE34), one seawater (IIE35) and two petroleum waters (IIE36 and IIE37). Each sample was filled in 250 ml Nalgene (high-density polyethylene, HDPE) bottles to distribute among the participating laboratories. Similarly, the CO₂ partial pressure of each sample was lower than the atmospheric partial pressure of CO₂ and samples were treated in the closed containers during preparation. Alteration due to the atmospheric CO₂ was therefore unlikely. The sample preparation and handling procedure was to insure the stability and homogeneity of the samples; however, no specific test was performed to verify them.

For chemical analysis of carbonic species, each laboratory was asked to report the preparation and calibration of standards, and analyses of the samples with two procedures: (a) regular laboratory procedure (or conventional procedure) and (b) forward and backward titration curves of each sample in tabular form. After the forward titration at pH<3 the titrand was left with magnetic stirring for 5 minutes for its CO₂ removal. This procedure for complete removal of CO₂ from the titration was used by Verma (2004) instead of bubbling with pure nitrogen or air with NaOH scrubber after the forward titration (Giggenbach and Goguel 1989, PNOC 2001). Then the backward titration was performed up to the original pH. Most of the laboratories ran the samples in duplicate and their repeated values were very close (good repeatability of individual laboratory).

From the titration curves the concentrations of carbonic species are calculated using the (1) Hydrologist's method, (2) Geochemist's method and (3) initial pH and total alkalinity method. Verma et al. (2014) reviewed the analytical procedures to optimize laboratory procedures for precise and accurate determination of carbonic species concentration in various types of waters. To apply the Hydrologist's and/or Geochemist's methods it is necessary to locate always two equivalence points (EP) for NaHCO₃EP and H₂CO₃EP, even for samples with pH lower than that of NaHCO₃EP. In addition, the backward titration curve with NaOH after complete removal of CO₂ is strictly necessary to decide for the applicability of the Hydrologist's method or the Geochemist's method. In cases where the complete analyses of species that contribute to the alkalinity are known, the initial pH and total alkalinity method is appropriate.

4. RESULTS AND DISCUSSION

Each laboratory reported the results using its own laboratory procedure and acid-base titration data in tabular form using a standardized procedure instruction for the acid-base titration. The organizer recalculated the reported parameters from the titration curves for all the participating laboratories and presented a comparative statistical evaluation of the results. Although all laboratories followed the general procedure for the acid-base titration curves, the participating laboratories were free to choose sample volume and concentration of acid-base standards according to their routine practice. Therefore, all titration datasets were converted first for the analysis of 25 ml sample volume. Some of the participating laboratories reported slightly different initial pH value of same sample in the datasets for the acid-base titration curves and the lab method. We considered the initial pH of acid-base titration curves as the sample pH in this work. Based on their origin, the water samples are classified into five groups:

4.1 Synthetic Samples

The synthetic water samples IIE26 to IIE29 were prepared with concentration higher than that of natural waters to examine the complete removal of CO₂ before the backward titration. Figure 2a shows the titration curves of sample IIE26, obtained from the reported tabular acid-base titration data of each laboratory. The effective amount was plotted on the x-axis to compare the titration curves of different laboratories. To handle the difference in the concentration of acid-base standards among different laboratories, a term "effective amount (meq)" is defined as $N \times V$, where N and V are the concentration (eq/l) and volume (ml) of acid-base added to the titrand, respectively. To plot together the forward and backward titration curves, the effective amount of acid-base in the titrand during the backward titration is considered as $V_a N_{acid} - V_b N_{base}$, where V_a (ml) is the total volume of acid of concentration (N_{acid} in eq/l) for the whole forward titration and V_b (ml) is the cumulative volume of base of concentration (N_{base} in eq/l) to the corresponding titration point. Since the initial pH of synthetic waters is higher than 10.5, the Hydrologist's method is applied for the comparative analysis of samples. These waters contain only the carbonic and OH⁻ alkalinity. The "dpH/dVol peak" method was used for locating the equivalence points; however, there is a certain displacement between the equivalence point and the peak in the dpH/dVol curve in some datasets (e.g., lab no. 3, 4 and 5 in Figure 2b). This is due to the fact that the number of titration points was not sufficiently large near the equivalence points.

Table 1 presents the results of the determination of carbonic species concentration by each laboratory conventional procedure. The carbonic alkalinity is calculated as $[\text{HCO}_3^-] + 2[\text{CO}_3^{2-}]$ from the reported concentration values of the individual laboratory method. For example, the carbonic alkalinity is higher than the total alkalinity (e.g., for lab no. 1 and 7) for sample IIE26, which is conceptually incorrect. Other possibility of this inconsistency is the higher analytical uncertainty in the results of laboratories 1 and 7. Sample IIE26 has only carbonic alkalinity as it contains only Na₂CO₃. Therefore, A must be equal to B. This is in reasonable agreement in all the datasets except for lab no. 7. However, the reported values for the individual laboratory method of some laboratories show high concentration of HCO₃⁻ (e.g., lab no. 4, 6 and 8) and OH⁻ (e.g. lab no. 5). Similarly, the titration curves of lab no. 7 show 16.4 mmol/l OH⁻. Theoretically, there is no appreciable concentration of OH⁻ and HCO₃⁻ in the sample.

There is only one laboratory that has the backward titration curve close to the theoretical backward titration curve (lab no. 8 in Figure 2a); the backward titration curves of other laboratories show considerable amount of CO₂ in the titrand. Thus it is recommended for future analyses to make sure the complete removal of CO₂ from the titrand by either bubbling with pure nitrogen or air with NaOH scrubber after the forward titration (PNOC 2001) or increase the magnetic stirring time before the backward titration (Verma 2004). The varying amount of CO₂ in the backward titration of the individual laboratory is probably a consequence of different rotation frequency and size of magnet used during the stirring of the titrand. A backward titration that is free from CO₂ is fundamental in defining the applicability of the methods for the determination of carbonic species concentration.

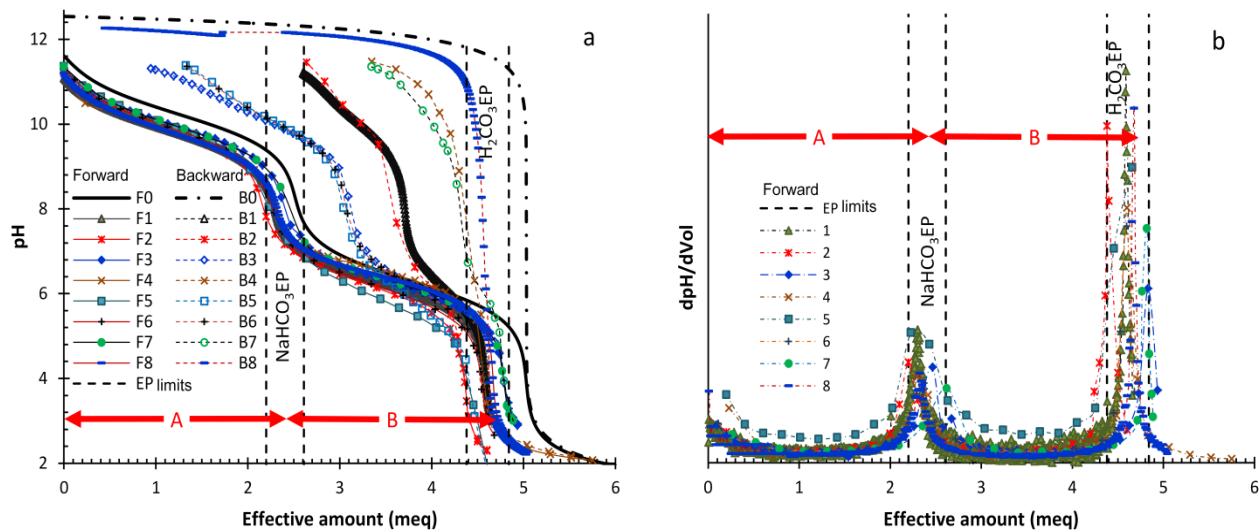


Figure 2: (a) Experimental and theoretical titration curves of synthetic sample IIE26 and (b) corresponding $d\text{pH}/d\text{Vol}$ curves for the forward titration curves. The vertical dashed lines define the range (i.e. maximum and minimum) of the locations of NaHCO_3EP and $\text{H}_2\text{CO}_3\text{EP}$. A and B are the averaged effective amounts (Verma et al. 2014).

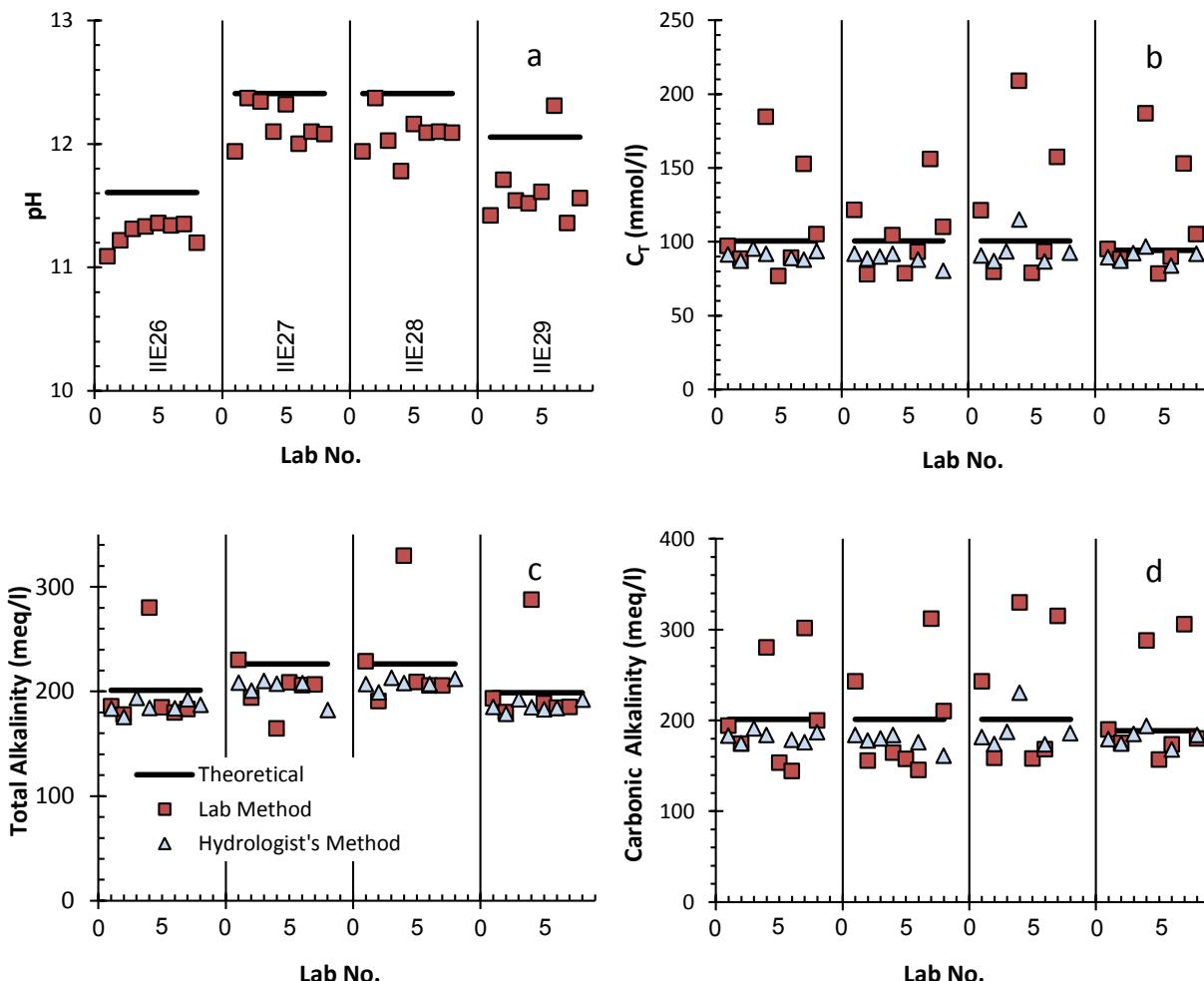


Figure 3: Behaviors of (a) pH, (b) C_T , (c) carbonic alkalinity and (d) total alkalinity of samples IIE26 to IIE29 (Verma et al. 2014).

Table 1. Concentration of carbonic species of synthetic samples (IIE26 to IIE29), analyzed according to individual laboratory procedure by the participating laboratories

Lab No.	Sample Vol. (ml)	Acid Conc. (eq/l)	Vol. added for total Alk. (ml)	Total Alk. (meq/l)	Concentration (mmol/l)				Calc. C _T	Calc. Carb. Alk (meq/l)
					pH	OH ⁻	HCO ₃ ⁻	CO ₃ ²⁻		
Sample IIE26 (Tot. Alk. = 201.2 meq/l, Carb. Alk. = 201.2 meq/l, C _T = 100.6 mmol/l, pH = 11.61, OH ⁻ = 0 mmol/l)										
1.	25.0	0.10	46.31	185.60	11.09			97.10	97.10	194.20
2.	25.0	1.60	2.75	177.78	11.22	3.39	2.74	85.83	88.56	174.39
3.					11.31					
4.	10.0	0.09	30.44	280.14*	11.33		89.02	95.56	184.58	280.14
5.	10.0	0.20	9.30	184.61	11.36	31.10		76.76	76.76	153.51
6.	50.0	0.02	456.55	179.88	11.34		34.21	54.92	89.13	144.05
7.	20.0	0.10	37.05	183.21	11.35		3.77	148.87	152.64	301.51
8.	50.0	0.05			11.20		10.00	95.00	105.00	200.00
Mean				182.22	11.28		27.95	93.43	113.40	206.83
S.D.				3.29	0.10		36.44	28.62	39.77	61.08
Sample IIE27 (Tot. Alk. = 226.2 meq/l, Carb. Alk. = 201.2 meq/l, C _T = 100.6 mmol/l, pH = 12.41, OH ⁻ = 25.0 mmol/l)										
1.	25.0	0.10	57.42	230.13	11.94			121.43	121.43	242.86
2.	25.0	1.60	3.00	194.02	12.37	38.32	0.20	77.75	77.95	155.70
3.					12.34					
4.	20.0	0.09	35.75	164.50	12.10		44.40	60.05	104.45	164.50
5.	10.0	0.20	10.50	208.43	12.32	51.21		78.74	78.74	157.48
6.	50.0	0.20	51.75	205.97	12.04		40.90	52.17	93.06	145.23
7.	10.0	0.10	20.90	206.70	12.10	57.09		155.97	155.97	311.95
8.	50.0	0.05			12.08		10.00	100.00	110.00	210.00
Mean				201.63	12.16	48.87	23.87	92.30	105.94	198.25
S.D.				21.63	0.16	9.60	22.09	36.53	27.23	61.21
Sample IIE28 (Tot. Alk. = 226.2 meq/l, Carb. Alk. = 201.2 meq/l, C _T = 100.6 mmol/l, pH = 12.41, OH ⁻ = 25.0 mmol/l)										
1.	25.0	0.10	57.05	228.67	11.94			121.37	121.37	242.74
2.	25.0	1.60	2.98	192.85	12.47	34.09	0.21	79.28	79.49	158.77
3.					12.03					
4.	10.0	0.09	35.83	329.74	11.78		88.04	120.87	208.91	329.77
5.	10.0	0.20	10.50	208.95	12.16	51.10		78.94	78.94	157.88
6.	50.0	0.20	51.70	205.77	12.09		18.49	74.92	93.42	168.34
7.	10.0	0.10	20.80	205.71	12.10	49.45		157.42	157.42	314.83
8.	50.0	0.05			12.09					
Mean				228.62	12.08	44.88	35.58	105.47	123.26	228.72
S.D.				50.87	0.20	9.38	46.34	33.20	51.53	79.22
Sample IIE29 (Tot. Alk. = 198.7 meq/l, Carb. Alk. = 188.7201.2 meq/l, C _T = 100.6 mmol/l, pH = 11.61, OH ⁻ = 0 mmol/l)										
1.	25.0	0.10	48.24	193.33	11.42			95.00	95.00	190.00
2.	25.0	1.60	2.82	180.04	11.71	7.10	1.32	86.81	88.13	174.95
3.					11.54					
4.	10.0	0.09	31.26	287.69	11.52		85.86	100.93	186.79	287.72
5.	10.0	0.20	9.50	188.58	11.61	31.10		78.41	78.41	156.82
6.	50.0	0.20	46.31	184.29	12.31		6.07	83.78	89.85	173.63
7.	10.0	0.10	18.70	184.94	11.36	3.64		152.88	152.88	305.77
8.	50.0	0.05			11.56		30.00	75.00	105.00	180.00
Mean				186.24	11.63	13.94	30.81	96.12	113.72	209.84
S.D.				4.99	0.30	14.96	38.79	26.60	40.35	60.40

* The underlined values were not considered in the calculation of mean and standard deviation.

The carbonic species concentrations for samples IIE27, IIE28, and IIE29 were also calculated by the Hydrologist's method. These waters have carbonic and OH⁻ alkalinities; we observed the same types of problems as discussed for sample IIE26. In future inter-laboratory comparison it is suggested to prepare synthetic samples with other types of alkalinity like boric and sulphide alkalinities.

The comparative behaviour of pH, total dissolved carbon (C_T), carbonic alkalinity and total alkalinity of the synthetic samples are shown in Figure 3. The conventional true values of each parameter are denoted by solid horizontal line for each sample. There is a good correlation between the values calculated from the tabular data of acid-base titration curves and the theoretical values. In contrast, the reported values based on the individual laboratory method for each sample have a high dispersion (i.e., large scatter). This suggests that there is a problem in the calculation procedure of the individual laboratory method for some of the participants.

4.2 Lake water

Figure 4 presents the titration curves of IIE30 (lake water), obtained from the tabular titration data of individual laboratory. The backward titration curves are close to the forward titration curves for most of the laboratories. This means that there are other types of alkalinity in the water sample and/or there is no complete removal of CO₂ after the forward titration. In the analyses of synthetic samples only partial removal of CO₂ after the forward titration has been observed. Thus it is difficult to present the correct evaluation of the backward titration of this sample.

Table 2 presents the carbonic species concentration of sample IIE30. It contains the values reported by the individual laboratory method and calculated values from the acid-base titration curves using the Hydrologist's and Geochemist's methods. Four laboratories reported the concentration of HCO_3^- and CO_3^{2-} , whereas other four laboratories reported only the concentration of HCO_3^- . The carbonic alkalinity is higher than the total alkalinity for lab. no. 4, 5 and 7. Theoretically, it is incorrect. The average value of pH of all laboratories is 8.29 ± 0.15 , which is quite close to the NaHCO_3EP . The acid-base titration curves of all laboratories show the existence of both equivalence points (i.e., NaHCO_3EP and $\text{H}_2\text{CO}_3\text{EP}$), except for lab no. 2 (Figure 4). Actually, lab no. 2 added a substantially large amount of acid at the beginning of titration (1 ml HCl 0.01 eq/l; pH changes from 8.55 to 7.62). Thus, we consider the location of NaHCO_3EP at pH=8.23 in this case.

The backward titration curve must always be above the forward titration curve. The backward titration curves of lab nos. 4 and 8 are lower than the corresponding forward titration curves. Both laboratories had added a considerable amount of effective acid during the forward titration (e.g., added effective amounts of 391 and 569 meq, respectively), whereas there is need of only 80 to 100 meq for locating both equivalence points. Thus there is probably a problem related to the high dilution of titrand in both laboratories.

The Hydrologist's method is appropriate for surface water samples. The calculated values using the acid-base titration method performed by all laboratories are in agreement (total alkalinity = 2.59 ± 0.15 meq/l, $C_T = 2.25 \pm 0.14$ mmol/l). If there is no complete removal of CO_2 before performing the backward titration, the Geochemist's method is not applicable. Therefore, the value of total dissolved CO_2 obtained by the Geochemist's method ($C_T = 1.78 \pm 0.46$ mmol/l) is lower than that of Hydrologist's method ($C_T = 2.25 \pm 0.14$ mmol/l). The partial removal of CO_2 during the backward titration also causes a high dispersion in the carbonic species concentration, calculated by the Geochemist's method.

4.3 Geothermal waters

The titration curves of all participating laboratories for geothermal water sample IIE31 are shown in Figure 5. The titration curves for samples IIE32 to IIE34 were similar to that of IIE31. These types of waters contain carbonic as well as considerable amounts of other alkalinites (e.g., boric and silicic alkalinites). Thus the Hydrologist's method is not suitable for the determination of carbonic species in these waters. The prerequisite of the Geochemist's method is the backward titration curve with complete removal of CO_2 . Due to the observations in the synthetic samples, a complete CO_2 removal cannot be ensured. Therefore, we applied the initial pH and total alkalinity method for comparison purposes.

Table 3 presents the carbonic species concentration of water samples IIE31 to IIE34, reported by the participating laboratories and calculated from the initial pH and total alkalinity method. The average values of pH of these samples are 7.35 ± 0.19 (IIE31), 7.11 ± 0.11 (IIE32), 7.38 ± 0.11 (IIE33), and 7.17 ± 0.26 (IIE34), respectively. It seems that all laboratories titrated the samples to the $\text{H}_2\text{CO}_3\text{EP}$ and converted the total alkalinity as the concentration of HCO_3^- . Laboratories nos. 1 and 2 reported the concentration of CO_3^{2-} as 0.42 and 0.01 mmol/l and pH as 7.30 and 7.60 for sample IIE31, respectively. Theoretically, there is no appreciable CO_3^{2-} at pH 7.30 or 7.60. Comparing the acid-base titration curves of all laboratories for sample IIE31, there was only one equivalence point ($\text{H}_2\text{CO}_3\text{EP}$).

As mentioned earlier the Hydrologist's method and Geochemist's method are based on the location of two equivalence points. For samples with an initial pH below the NaHCO_3EP (pH ≈ 8.2) there is need to first titrate these samples with NaOH (or to add a certain amount of NaOH) to reach a pH above the equivalence points. It is recommended here to add a NaOH standard solution until a pH of ~ 9.0 is reached and then titrate the titrand with the acid standard to locate both equivalence points.

To apply the initial pH and total alkalinity method it is necessary to locate the $\text{H}_2\text{CO}_3\text{EP}$ in the forward acid-titration curves of each laboratory (Figure 5). The total alkalinity is calculated by dividing the effective amount for the $\text{H}_2\text{CO}_3\text{EP}$ by the sample volume. Knowing the initial pH and the concentration of boron and SiO_2 , the non-carbonic alkalinity is calculated in Table 3. The carbonic alkalinity is the total alkalinity minus the non-carbonic alkalinity. From the carbonic alkalinity and initial pH the concentration of individual carbonic species is calculated. The samples have 1% (IIE31), 44% (IIE32), 32% (IIE33), and 86% (IIE34) of non-carbonic alkalinity. Thus there is a considerable error in the determination of carbonic species by using the method of converting total alkalinity to HCO_3^- as performed by the participating laboratories.

4.4 Seawater

The titration curves of seawater sample (IIE35) are presented in Figure 6. The lab no. 3, 4, and 8 had problems in their titration curves and were not considered in the calculations. The values reported by the individual laboratory and calculated from the acid-base titration curves using the Hydrologist's method are in good agreement. This suggests that the analytical procedure of most of the participating laboratories is similar to the Hydrologist's method. The Hydrologist's method requires both NaHCO_3EP and $\text{H}_2\text{CO}_3\text{EP}$ equivalence points. In case of seawater, the initial pH is very close to the pH of NaHCO_3EP and the location of $\text{H}_2\text{CO}_3\text{EP}$ is sufficient in this specific case.

The average values of total alkalinity (also equal to carbonic alkalinity) calculated by individual laboratory procedure (2.22 ± 0.63 meq/l) and by Hydrologist's method (2.50 ± 0.05 meq/l) are in agreement with the averaged total alkalinity (2.31 meq/l) of surface ocean waters obtained under the long term monitoring project (HOT-DOGS 2012).

4.5 Petroleum waters

The acid-base titration curves for petroleum sample IIE36 are shown in Figure 7. The forward and backward titration curves are similar and there is no well-defined inflection point even for $\text{H}_2\text{CO}_3\text{EP}$. The same behaviour of petroleum waters even with complete removal of CO_2 from the titrand before the backward titration was observed earlier (Verma et al. 2010). Actually, if the backward titration overlaps the forward titration curve, there is no carbonic alkalinity and consequently no carbonic species concentration in the petroleum waters. There is still need to perform a study with measuring all relevant species in petroleum waters.

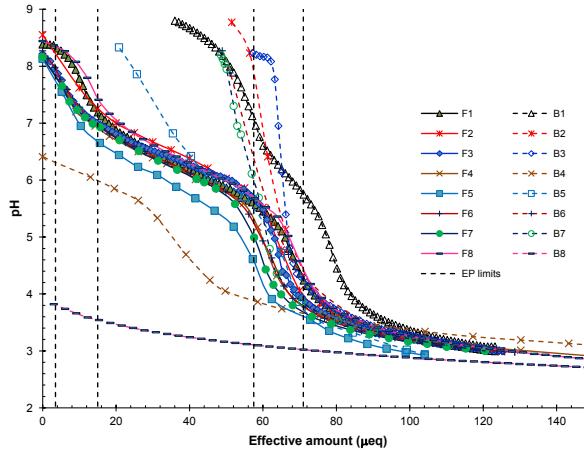


Figure 4: Forward and backward titration curves of lake water sample IIE30. The parameters are defined in Figure 1.

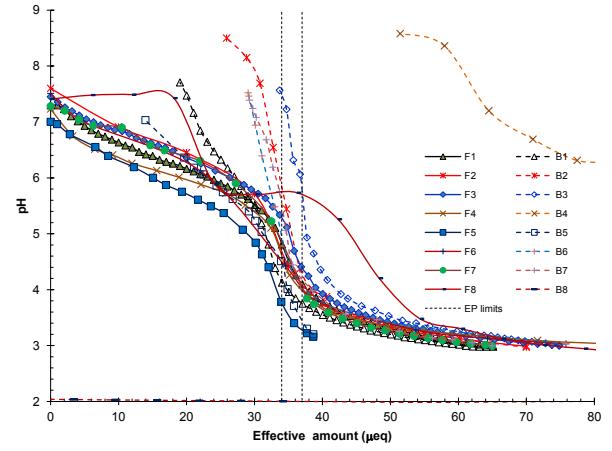


Figure 5: Forward and backward titration curves of geothermal water sample IIE31. The parameters are defined in Figure 1.

Table 2. Concentration of carbonic species of lake water sample (IIE30) analyzed with individual laboratory method and with the Hydrologist's and Geochemist's methods from the titration curves performed by the participating laboratories.

Lab No.	Parameters ^ξ (μeq)				pH	Total Alk. (meq/l)	Carb. Alk. (meq/l)	Concentration (mmol/l)		
	A'	B'	C'	D'				HCO ₃ ⁻	CO ₃ ²⁻	C _T
Individual laboratory analysis										
1.					8.38	2.77	2.77	1.93	0.42	2.35
2.					8.55	2.57	2.56	2.56	0.09	2.47
3.					8.22	2.40	2.40	2.40		2.40
4.					8.13	2.17	2.71	2.71	0.17	2.55
5.					8.17	3.26	3.27	2.27	0.60	2.67
6.					8.27	0.14 ^a	2.38	2.38		2.38
7.					8.16	1.37	2.24	2.24		2.24
8.					8.44		2.90	2.90		2.90
Mean					8.29	2.42	2.65	2.42	0.32	2.50
S.D.					0.15	0.63	0.33	0.30	0.23	0.21
Hydrologist's method										
1.	11.5	57				2.74	2.74	1.82	0.46	2.28
2.	3.5	64				2.70	2.70	2.42	0.14	2.56
3.	6	58				2.56	2.56	2.08	0.24	2.32
4.	9.1	54.3				2.54	2.54	1.81	0.36	2.17
5.	8.5	53				2.46	2.46	1.78	0.34	2.12
6.	7	54.5				2.46	2.46	1.9	0.28	2.18
7.	7	54				2.44	2.44	1.88	0.28	2.16
8.	15	56				2.84	2.84	1.64	0.6	2.24
Mean						2.59	2.59	1.92	0.34	2.25
S.D.						0.15	0.15	0.24	0.14	0.14
Geochemist's method										
1.	11.5	57	14	8		2.74	1.86	1.58	0.14	1.72
2.	3.5	64	11.09	2.21		2.70	2.17	2.06	0.05	2.12
3.	6	58	0.8	5.9		2.56	2.29	2.28	0.00	2.29
4.	9.1	54.3				2.54				
5.	8.5	53	29.3	9.5		2.46	0.91	0.99	-0.04	0.95
6.	7	54.5	8.7	4		2.46	1.95	1.71	0.12	1.83
7.	7	54	9.2	3		2.44	1.95	1.63	0.16	1.79
8.	15	56				2.84				
Mean						2.59	1.86	1.71	0.07	1.68
S.D.						0.15	0.49	0.45	0.08	0.44

^aThe underlined value was not considered in the calculation of mean and standard deviation.

^ξParameters A', B', C' and D' are defined by Verma et al. (2014).

Table 3. Concentration of carbonic species of geothermal water samples (IIE31 to IIE34) analyzed with individual laboratory method, and with the initial pH and total alkalinity method using the titration curves performed by the participating laboratories.

Lab No.	Laboratory method							Initial pH and total alkalinity method using titration curves of individual laboratory								
	Sample Vol (ml)	Acid (eq/l)	Acid Vol (ml)	pH	Alkalinity (meq/l)		Concentration (mmol/l)	Effective amount (μeq)	Alkalinity (meq/l)			Concentration (mmol/l)				
					Total Alk	Car. Alk.			HCO ₃ ⁻	CO ₃ ²⁻	C _T	Total Alk.	Non Carb. Alk.	Carb. Alk.	H ₂ CO ₃	HCO ₃ ⁻
Sample IIE31 (SiO ₂ = 2.16 mmol/l, B=0.31 mmol/l)																
1	25	0.10	0.69	7.30	2.77	1.40	1.16	0.12	1.28	34.8	1.40	1.01E-2	1.38	0.14	1.38	1.52
2	25	0.16	0.20	7.60	1.27	1.27	1.26	0.01	1.27	35.5	1.42	2.01E-2	1.40	0.07	1.39	1.47
3	50	0.02		7.45		1.42	1.42		1.42	35.9	1.44	1.42E-2	1.42	0.10	1.42	1.52
4	20	0.01	2.76	7.23	1.44	1.44	1.44		1.44	35.2	1.41	8.62E-3	1.40	0.16	1.40	1.56
5	10	0.02	0.80	7.00	1.51	1.55	1.55		1.55	34.0	1.36	5.09E-3	1.35	0.27	1.35	1.62
6	50	0.02	4.80	7.51	1.89	1.55	1.55		1.55	37.0	1.48	1.64E-2	1.46	0.09	1.46	1.55
7	25	0.02	2.00	7.28	1.58	1.30	1.30		1.30	37.0	1.48	9.67E-3	1.47	0.15	1.47	1.62
8	50	0.05		7.41		2.20	2.20		2.20							
Mean					7.35	1.74	1.42	0.07	1.40	35.63	1.43	0.01	1.41	0.14	1.41	1.55
S.D.					0.19	0.54	0.11	0.145	0.12	1.11	0.04	0.01	0.04	0.07	0.04	0.06
Sample IIE32 (SiO ₂ = 15.76 mmol/l, B=51.64 mmol/l)																
1	25	0.10	0.25	7.15	1.01	1.00	1.00		1.00	24.1	0.96	0.46	0.50	0.07	0.50	0.58
2	25	0.16	0.13	7.30	0.84	0.83	0.83		0.83	23.2	0.93	0.65	0.28	0.03	0.28	0.31
3	50	0.02		7.16						24.4	0.98	0.47	0.50	0.07	0.50	0.57
4	20	0.01	1.91	7.03	0.79	0.99	0.99		0.99	25.0	1.00	0.35	0.65	0.12	0.65	0.77
5	25	0.02	1.45	6.95	1.10	1.11	1.11		1.11	28.0	1.12	0.29	0.83	0.19	0.83	1.01
6	50	0.02	18.68	7.03	0.37	0.88	0.88		0.88	24.1	0.96	0.35	0.61	0.11	0.61	0.73
7	25	0.02	1.45	7.08	0.57	0.93	0.93		0.93	23.0	0.92	0.39	0.53	0.09	0.53	0.62
8	50	0.05		7.16		1.80	1.80		1.80	24.0	0.96	0.47	0.49	0.07	0.49	0.56
Mean					7.11	0.78	0.96	0.96	0.96	23.97	0.96	0.43	0.55	0.09	0.55	0.64
S.D.					0.11	0.27	0.10	0.10	0.10	0.68	0.03	0.11	0.16	0.05	0.16	0.20
Sample IIE33 (SiO ₂ = 17.46 mmol/l, B=33.76 mmol/l)																
1	25	0.10	0.29	7.38	1.16	1.16	1.16		1.12	28.0	1.12	0.53	0.59	0.05	0.58	0.63
2	25	0.16	0.13	7.55	0.87	0.87	0.87		1.05	26.3	1.05	0.78	0.27	0.01	0.27	0.28
3	50	0.02		7.46					1.14	28.5	1.14	0.64	0.50	0.03	0.50	0.53
4	20	0.01	2.23	7.23	0.93	1.16	1.16		1.16	29.0	1.16	0.38	0.78	0.09	0.78	0.87
5	10	0.02	0.70	7.23	1.32	1.33	1.33		1.08	27.0	1.08	0.38	0.70	0.08	0.70	0.78
6	50	0.02	18.72	7.43	0.37	0.57	0.57		1.07	26.7	1.07	0.60	0.47	0.03	0.47	0.50
7	25	0.02	1.60	7.33	0.63	1.03	1.03		1.09	27.2	1.09	0.48	0.61	0.06	0.61	0.67
8	50	0.05		7.41		1.00	1.00		1.21	30.3	1.21	0.57	0.64	0.05	0.64	0.69
Mean					7.38	0.98	1.09	1.09	1.02	27.88	1.12	0.55	0.57	0.05	0.57	0.62
S.D.					0.11	0.27	0.16	0.16	0.25	1.34	0.05	0.14	0.16	0.03	0.16	0.18
Sample IIE34 (SiO ₂ = 15.54 mmol/l, B=27.38 mmol/l)																
1	25	0.10	0.09	7.30	0.36	0.35	0.19	0.08	0.27	8.2	0.33	0.36	-0.04	0.00	-0.04	-0.04

2	25	0.16	0.04	7.44	0.25	0.25	0.25	7.9	0.32	0.50	-0.19	-0.01	-0.18	-0.20	
3	50	0.02		7.26		<u>6.59*</u>	<u>6.59</u>	6.59	8.2	0.33	0.34	-0.01	0.00	-0.01	-0.01
4	20	0.01	0.65	7.27	0.27	0.34	0.34	0.34	9.4	0.38	0.34	0.04	0.00	0.04	0.04
5	25	0.02	0.50	6.64	0.39	0.39	0.39	0.39	7.1	0.28	0.08	0.20	0.09	0.20	0.30
6	50	0.02	5.77	6.96	<u>0.11</u>	0.38	0.38	0.38	8.8	0.35	0.17	0.18	0.04	0.18	0.22
7	25	0.02	0.50	7.16	0.21	0.34	0.34	0.34	8.3	0.33	0.27	0.07	0.01	0.07	0.08
8	50	0.05		7.37		<u>3.40</u>	<u>3.40</u>	3.40	15.0	0.60	0.43	0.17	0.01	0.17	0.19
Mean				7.18	0.30	0.34	0.31	0.08	1.50	0.33	0.31	0.05	0.02	0.05	0.07
S.D.				0.26	0.08	0.05	0.08	2.32	0.72	0.03	0.14	0.13	0.03	0.13	0.16

* The underlined values were not considered in the calculation of mean and standard deviation.

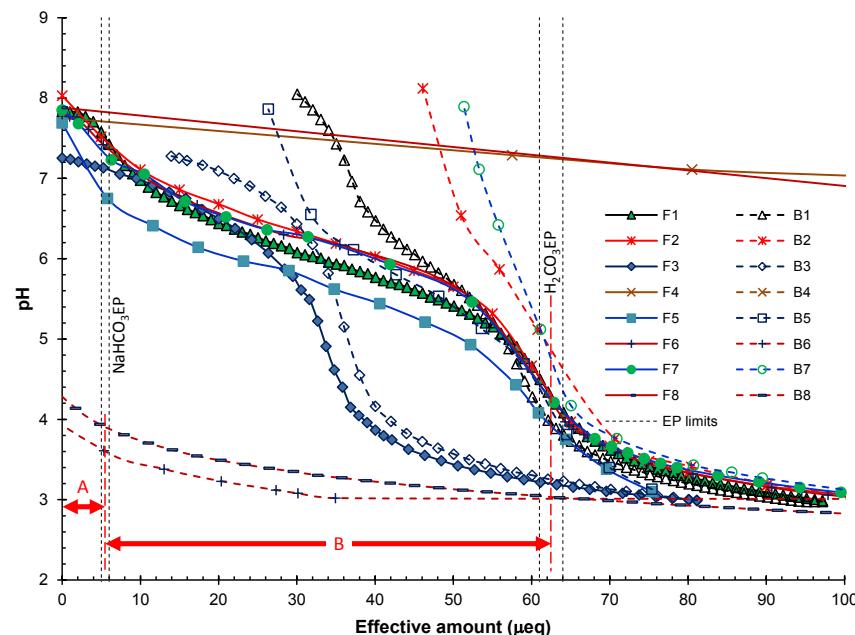


Figure 6: Forward and backward titration curves of seawater sample IIE35. The parameters are defined in Figure 1.

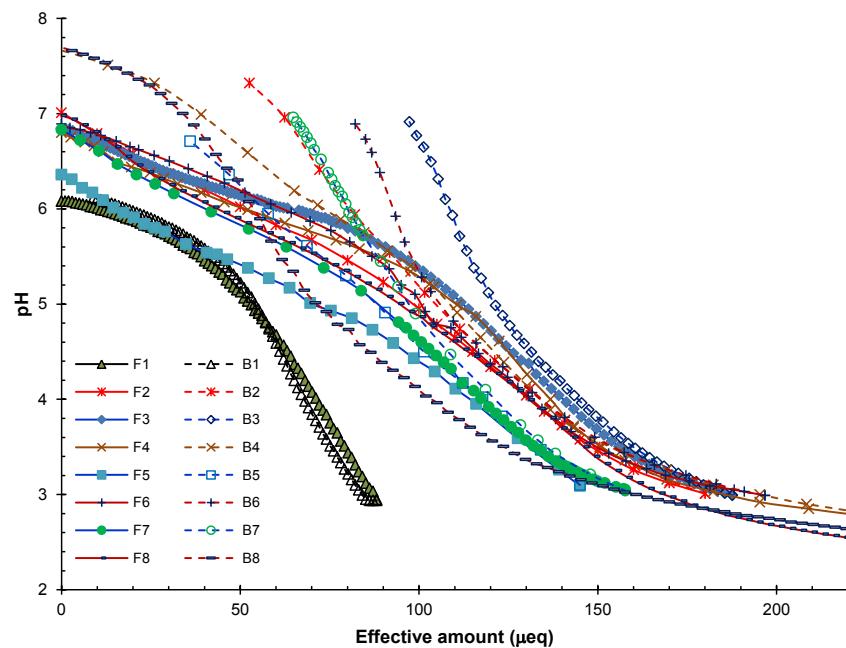


Figure 7: Forward and backward titration curves of petroleum water sample IIE36. The parameters are defined in Figure 1.

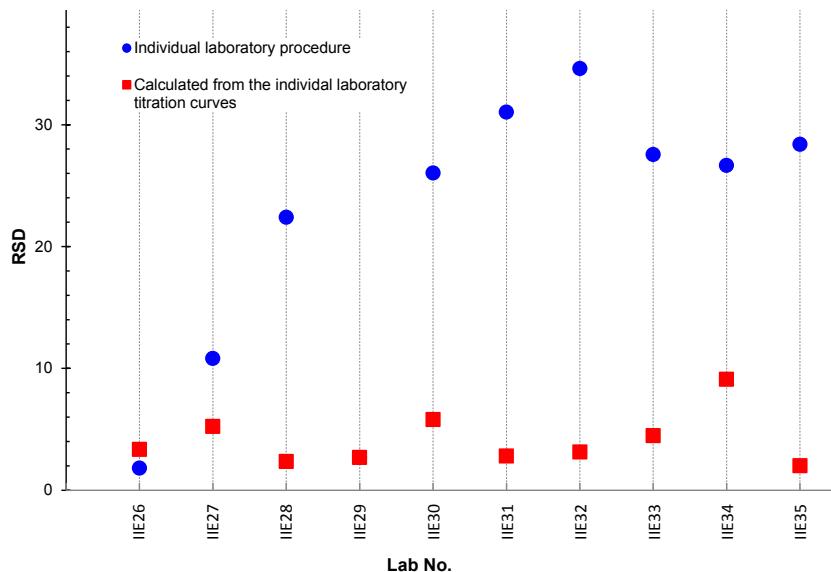


Figure 8: Comparison of analytical uncertainty in the total alkalinity, calculated from the reported results and from the titration curves of individual laboratory for each sample.

4.6 Measure data quality assessment

A comparison of RSD in the total alkalinity, calculated from the reported results and from the titration curves of individual laboratory for each sample was performed (Figure 8). The total alkalinity calculation method used here depends on the water type: the Hydrologist's method for synthetic samples (IIE26 to IIE29), lake water (IIE30) and seawater (IIE35) and the initial pH and total alkalinity method for geothermal waters (IIE31 to IIE34). The forward and backward titration curves for petroleum waters are overlapping; none of methods can be applied for the calculation of carbonic species concentration. If the concentrations of all the dissolved species are known, the *initial pH and total alkalinity method* can be applied with certain limitations.

The uncertainty for each sample obtained from the tabular titration curves is typically less than 10%, a value that is reasonable. In contrast, the uncertainties derived from individual laboratory procedures show systematic larger values between 10% and 35%, except for samples IIE26 (1.8%) and IIE29 (2.7%). This highlights the need in future inter-laboratory comparison of carbonic species distribution to perform a more detailed revision of the calculation procedures of individual laboratories.

4. CONCLUSIONS

The inter-laboratory comparison study for the determination of carbonic species concentration in twelve natural waters (four synthetic waters, one lake water, four geothermal waters, one seawater, and two petroleum waters), conducted among eight worldwide laboratories is summarized as follows:

- The acid-base titration procedure is the only reliable tool for the determination of carbonic species concentration in natural waters. It is classified in three methods: (1) *Hydrologist's method*, (2) *Geochemist's method*, and (3) *initial pH and total alkalinity method*.
- The back titration curve is fundamental in defining the applicability of *Hydrologist's* and/or *Geochemist's* methods.
- The *Hydrologist's* and/or *Geochemist's* methods always demand the location of two equivalence points (NaHCO_3EP and $\text{H}_2\text{CO}_3\text{EP}$); whereas *initial pH and total alkalinity method* is based on the location of only $\text{H}_2\text{CO}_3\text{EP}$.
- The addition/liberation of CO_2 from the titrand before the location of NaHCO_3EP may cause analytical uncertainty for the *Hydrologist's* and/or *Geochemist's* methods; whereas it has no effect in the *initial pH and total alkalinity method*.
- The analytical procedures of all the participating laboratories are close the *Hydrologist's method*. Thus their procedures are not applicable for geothermal and petroleum waters.
- High dispersion among the results obtained by the individual laboratory analytical procedure is observed. In contrast, the tabular titration curves performed by following a standardized protocol provide consistent results. The concentrations of carbonic species obtained by the latter titration curves are in good agreement with the conventional true values for synthetic waters. Consequently, a revision of calculation procedure of the laboratories is strongly advised.
- Finally, this study showed the need to measure all the chemical species in petroleum waters in order to define an appropriate titration procedure for the determination of carbonic species in this type of water.

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