

Recent Geological, Geochemical and Geophysical Surveys of the Roseau Valley, High-Temperature Geothermal Field in Dominica, West Indies

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ABSTRACT

The Roseau Valley in Dominica (West Indies) is characterized by recent volcanic activity (less than 50,000 years BP) with several eruptive centres and abundant surface manifestations, which are observed in two main spots distant of 4 kilometers: Wotten Waven in the Roseau Valley and the adjacent Boiling Lake –Valley of Desolation area. Several types of thermal manifestations have been recorded, including warm springs, hot springs, fumaroles and steam vents, cold gas discharges, solfatares, fossil alteration areas and phreatic craters. Some of the hot springs discharge Na-Cl and Ca-rich Na-Cl waters, which point out the existence of a high temperature geothermal reservoir. These waters might have a common origin, but would have evolved separately during their ascent to the surface. They have experienced water-rock interactions at high temperature as indicated by their oxygen-18 shift. The Na-Cl waters discharged in the Roseau Valley area have deep equilibrium temperature around 200-225°C. Ca-rich Na-Cl waters are discharged in the Valley of Desolation and the vicinity of the Boiling Lake. Their higher contents of calcium and chloride could be indicative of a strong degassing before they reach the surface. Na-K geothermometers indicate higher deep equilibrium temperatures (up to 300°C).

A geophysical (magnetotelluric) survey has been carried out with 32 MT soundings. Unfortunately, the natural MT signal was very low and the depth of penetration was limited (400 to 1000 m). However, most of the soundings record a shallow, highly conductive layer (1-7 ohm.m) with a thickness of 250-350 m ascribed to an argillic cap-rock. A few soundings also record an underlying, less conductive layer (15 to 80 ohm.m). This resistivity pattern is commonly observed in many high temperature geothermal reservoirs. In addition, the MT soundings support the hypothesis of a connection at the reservoir depth between Wotten Waven and the Boiling Lake –Valley of Desolation area.

The conceptual model which best fits in with the geological, geochemical and geophysical data is a deep Na-Cl reservoir extending below the Micotrin recent lava dome. Its temperature is estimated to be in the range of 200-300°C. Lateral fluid outflows develop to the southeast (Boiling Lake-Valley of Desolation) and to the southwest (Roseau Valley). The total size of the reservoir where high temperature fluids are thought to be encountered is estimated to be about 18 km². A preliminary assessment of the capacity of the western half of the reservoir has been done with a preliminary figure of 100-120 MWe. This is in accordance with the average power density of 10 MW per km² observed for high temperature geothermal reservoir worldwide.

1. INTRODUCTION

Dominica is located in the central Lesser Antilles between the two French islands Guadeloupe and Martinique. It is about 45 km long by 25 km wide and its area is about 800 km². Its morphology shows a N-S trending axial ridge formed by several distinct volcanic complexes and covered by dense tropical forests. Geothermal potential of Dominica has been first reported during an UNDP visit (Barnea et al., 1969). First detailed exploration has been done by the French Geological Survey (BRGM) between 1982 and 1984. It included geology, geochemistry and geophysical (gravimetric, resistivity, magnetotelluric) methods (BRGM, 1984 and 1985). Two areas of interest had been investigated: Wotten Waven and Soufriere regions. Wotten Waven is located in the Roseau River valley about 8 km ENE of the capital Roseau (Figure 1). Soufriere is located at the southern tip of the island. Later, Geotermica Italiana (1992) carried out an assessment of geothermal resources in the Eastern Caribbean, funded by UN-DTCD and CARICOM. Their report updated previous data gained by BRGM in 1982-84 and provided some conclusions on Dominica geothermal potential and feasibility of power transmission via submarine cable systems to neighbour islands.

More recently, a new exploration survey including field geology and fluid geochemistry was done in the region of Wotten Waven in the frame of the Eastern Caribbean Geothermal Development Programme "Geo-Caraïbes" funded by the OAS (Organisation of American States). This survey updated previous data and provided a preliminary conceptual model of the reservoir (Lasne and Traineau, 2005). In 2008, another programme called "Geothermal Energy in Caribbean Islands" or "Géothermie Caraïbes" was initiated by the E.U., the Commonwealth of Dominica and France in the frame of the European INTERREG III-B Programme "Espaces Caraïbes". It supported a prefeasibility study of the development of the Wotten Waven geothermal field, which is better named "the Roseau Valley geothermal field" because it extends widely outside the area of Wotten Waven village (Figure 1). It included additional field surveys in geology and fluid geochemistry (Traineau and Lasne, 2008; Traineau et al., 2008), and a new MT survey done by BRGM (Baltassat et al., 2008).

This paper intends to present the results obtained during the 2005 survey and the 2008 prefeasibility study. Based on these surveys, the Government of Dominica drilled five deep exploration wells between 2012 and 2014. Preliminary results, which are not in the scope of this paper, would have confirmed the existence of a high temperature geothermal reservoir suitable for power generation.

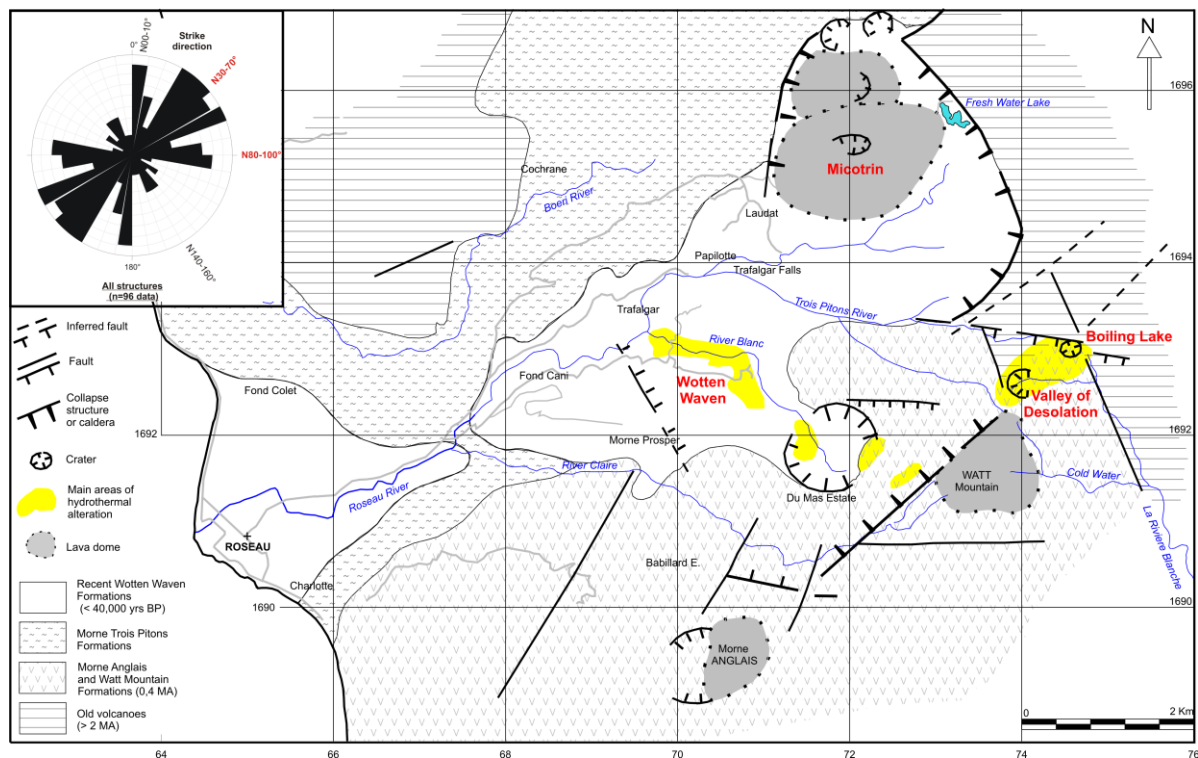


Figure 1: Simplified geological and structural map of the Roseau Valley [adapted from BRGM (1984) and Lasne and Traineau (2005)]. Rose diagram of strike of structures measured on the field (from Lasne and Traineau, 2005).

2. GEOLOGY

The Roseau Valley is characterized by a recent volcanic activity (less than 50,000 years BP; BRGM, 1985b) with several eruptive centres located in its northern and eastern margins (Figure 1). The most important event was the emission of the Roseau Ignimbrite which ended with the building of the Micotrin (or Morne Macaque) lava domes about 26,000-29,000 years ago (BRGM, 1985b). Recent radiocarbon ages provide evidence of a magmatic and chronological link between the Roseau Ignimbrite, Micotrin and Trois Pitons lava domes (BRGM, 1985b). They would all derive from the draining of a large andesitic to dacitic magmatic chamber. There is debate on the occurrence and location of a caldera, associated with the emission of the Roseau Ignimbrite Complex. This caldera might have been obliterated by younger pyroclastic deposits and lava domes. Previous authors proposed a large semi-circular volcanic depression opened to the Caribbean Sea (Robson and Tomblin, 1966), or a gravity-slide structure (Roobol et al., 1983). BRGM (1984) proposed a large collapse structure centred on the Wotten Waven depression. Lasne and Traineau (2005) did not find any convincing data for supporting a large caldera structure as proposed by previous authors. They considered that the proposed link between Morne Trois Pitons activity and Roseau Ignimbrite/Micotrin activity does not require or support a large caldera centred on the Wotten Waven depression. They only outlined limits of a 2-km diameter caldera filled with the two coalescent domes of Micotrin and its extension southwards, which could be a remnant of a large ignimbrite caldera (Figure 1).

Other small magmatic eruptions have been recorded in the Roseau Valley area. One would be associated with a sector collapse caused by a magmatic intrusion in Du Mas Estate area about 1,300 years ago (Figure 1), as proposed by BRGM (1984). Another corresponds to the emission of a dacitic pumiceous horizon 1,030 years ago from an unknown source. The circular-shape depression, including the Boiling Lake and the Valley of Desolation could be a candidate. This recent magmatic activity is related to differentiated magmas (acid andesite, dacite) with high viscosity, which promotes trapping and cooling of magma intrusions at shallow depths.

There are two conspicuous fault-line valleys trending NE-SW and NNW-SSE in the vicinity of the Boiling Lake and Valley of Desolation (Figure 1). The NNW-SSE trend is parallel to the alignment of the recent Trois Pitons and Micotrin lava domes. The high River Claire valley southwest of Watt Mountain is also considered as a NE-SW trending, fault-line valley. Lasne and Traineau (2005) pointed out significant fracturing in the Roseau Valley basement with several fracture sets (Figure 1). The main sets strike NE-SW, E-W and N-S. They are well correlated with the main inferred faults mapped in and around the Roseau Valley. The NE-SW fracture set preferentially controls shallow geothermal fluid circulations in the River Blanc valley. The Wotten Waven area is characterized with active seismicity as shown by a recent seismic swarm recorded in 1998-99 (Young, 2005) which promotes fracturing. This is supported by the occurrence of fractures in the recent debris flow deposits outcropping in River Blanc valley and dated at 1,3 ky.

3. SURFACE MANIFESTATIONS MAPPING

The most striking feature of the Roseau Valley geothermal field is the abundance of surface manifestations, which are recorded in two main spots distant of 4 kilometres: the River Blanc valley near the Wotten Waven Village (Sulphur Springs), and the Boiling Lake –Valley of Desolation area on the other side of the axial ridge (Figure 1 and Figure 2). Lasne and Traineau (2005) provided a

comprehensive map of the surface manifestations classified into eight types: warm spring, hot spring, mineralized-fluid hot spring, fumaroles, cold gas discharge, solfatare and steam vent, fossil alteration area, phreatic crater.

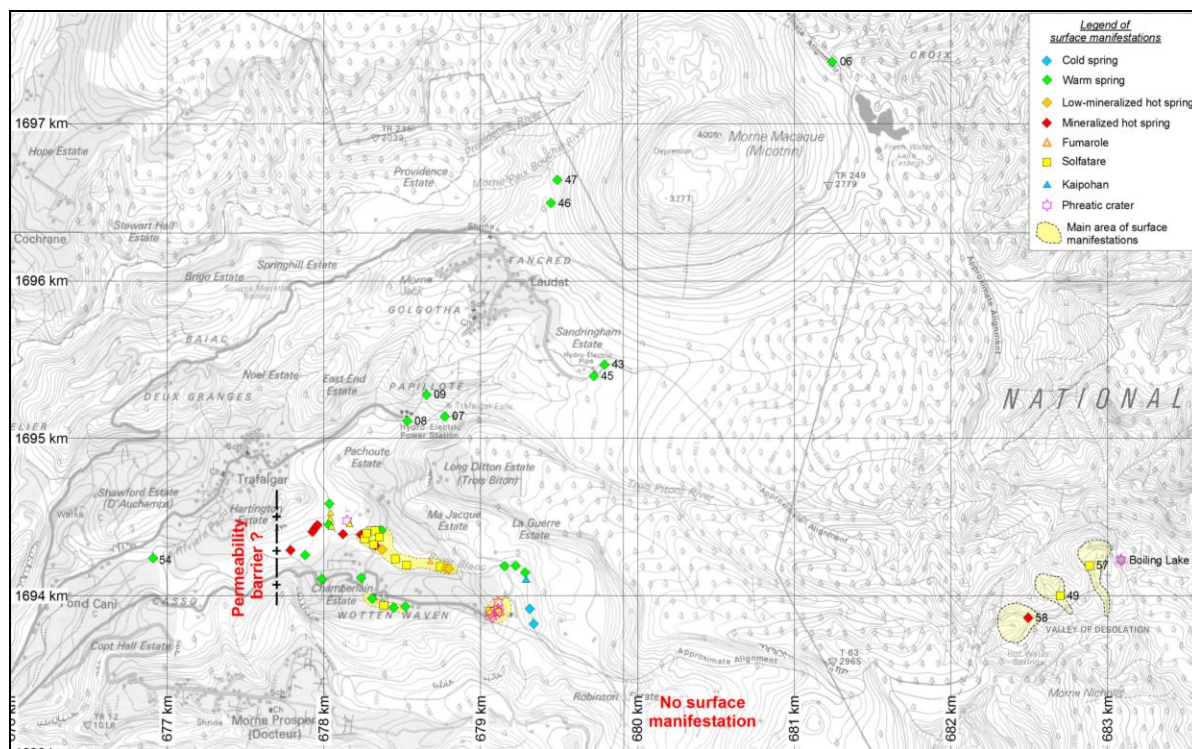


Figure 2 : Type and location of the surface manifestations of the Roseau valley geothermal field compiled by Traineau and Lasne (2008). Coordinates in UTM WGS84. Elevation in feet.

In River Blanc Valley, there is a clear zoning of surface manifestations with increasing elevation between 200 to 400 m.a.s.l. (Figure 3). In the lower section of River Blanc (downstream the bridge) and to some extent in the Roseau River and Trois Pitons River, there are several hot springs delivering Na-Cl waters and precipitating amorphous silica. Such hydrothermal surface manifestations are often considered to represent the distal part of a lateral fluid outflow coming from a deep, high temperature, sodium chloride reservoir. In the middle and upper sections of the River Blanc (upstream the bridge), in the River Camelia, there are abundant steam vents, solfatares and springs delivering hot and warm, low-mineralized fluids. Hydrothermal alteration is extensive with deposition of native sulphur, alunite, sulphates, Fe-sulphides, silica, clay material. It is ascribed to the advanced argilic alteration type. The vicinity of the old Wotten Waven lodge is characterized with the occurrence of several small-size phreatic vents associated with a low-intensity or fossil fumarolic activity. Several areas of cold gas discharges have been observed.

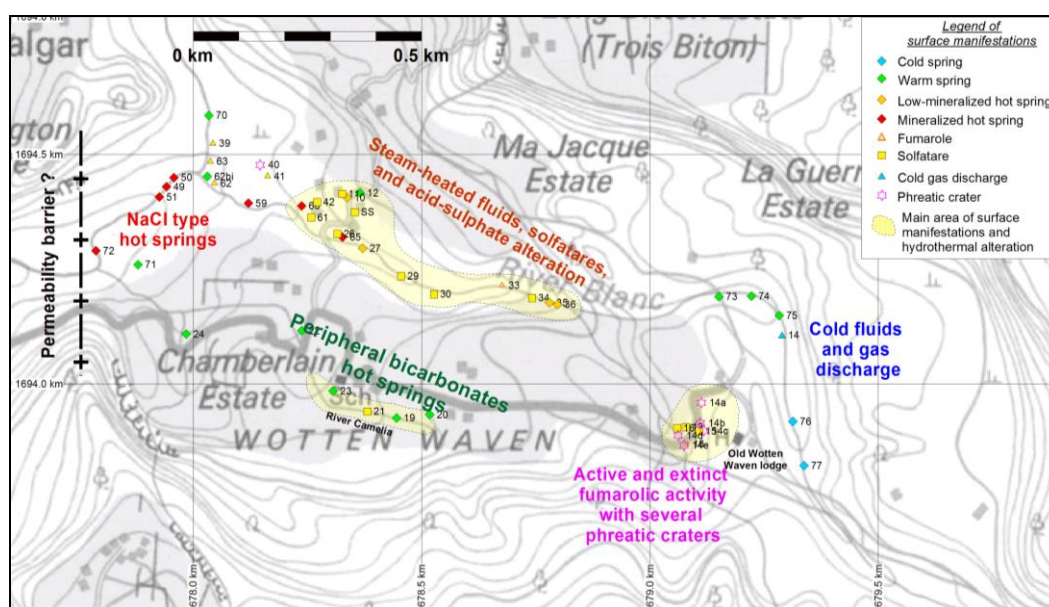


Figure 3 : Enlargement of Figure 2 showing a clear zoning of surface manifestations with increasing elevation in the River Blanc Valley from 200 to 400 masl (from Traineau and Lasne, 2008).



Figure 4: The phreatic crater of Boiling Lake (left) and the Valley of Desolation (right) which was the site of phreatic explosions in 1863 and 1880.

The Boiling Lake –Valley of Desolation area represents another main spot of surface manifestations (Figure 4). Its most prominent feature is the Boiling Lake, considered as a recent but undated phreatic crater. The Valley of Desolation area has been the site of phreatic explosions in 1863 and 1880. Today it is characterized mainly by solfataric activity, with several steam vents and hot springs. Some of them discharge mineralized fluids, which derive from a high-temperature reservoir. The Boiling Lake – Valley of Desolation hydrothermal area is located at an elevation of 700-800 metres within massive lavas belonging to Watt Mountain. A structural control is considered for Boiling Lake with the proximity of a NNW-SSE trending fault.

Other surface manifestations are scattered in the Papillote-Trafalgar Falls region and on the flanks of the Micotrin lava dome at elevation of 740-780 m.a.s.l. (Sandringham Estate, Fresh Water Lake, Paix Bouche River; Figure 2). They are warm springs discharging low-mineralized sodium-bicarbonate fluids. Warm springs on the flanks of the Micotrin lava domes are important because they point out the existence of shallow bicarbonate aquifers which are often considered as peripheral waters developed above deep high-temperature reservoirs.

The lack of thermal manifestation in the Roseau valley downstream of Wotten Waven (except one occurrence of warm fluid and gas discharging in the Roseau River close to its junction with La Riviere Padu) is also striking. This suggests the existence of a hidden permeability barrier as mentioned on Figure 2, which could be either a fault zone or a caldera border. The high valley of River Blanc (Robinson Estate, Du Mas Estate) and the River Clair valley are also lacking thermal manifestation (Figure 2). Also there is no known surface manifestation east of the Boiling Lake and Valley of Desolation area. This puts constraints in the location of the deep geothermal reservoir which should be developed within a triangle defined by Micotrin northwards, Boiling Lake/Valley of Desolation eastwards and Wotten Waven south-westwards (Figure 2).

3. FLUID GEOCHEMISTRY

About 35 warm and hot springs have been recorded in the Wotten Waven area and the Boiling Lake - Valley of Desolation area and analysed by BRGM (1985), Lasne and Traineau (2005) and Traineau and Lasne (2008). The most representative fluid samples are given in Table 1. They have been plotted in the Cl-SO₄-HCO₃ triangular diagram of Giggenbach (1988; Figure 5). They are without any doubt indicative of the existence of a high temperature geothermal reservoir with neutral sodium chloride waters, bicarbonate waters (Ca-Na-HCO₃ type and Na-HCO₃-SO₄ type) and acid-sulphate waters. In addition, Ca-rich sodium chloride waters have been identified in the Boiling Lake and Valley of Desolation area.

Table 1 : Selected chemical and isotopic data of Na-Cl waters and Ca-rich Na-Cl waters discharged by the thermal manifestations of Wotten Waven and Boiling-Valley of Desolation areas in Dominica [from BRGM (1985), Lasne and Traineau (2005), Traineau and Lasne (2008)].

Site	date	label	T (°C)	pH	Cl mg/l	HCO ₃ mg/l	SO ₄ mg/l	SiO ₂ mg/l	Na mg/l	K mg/l	Ca mg/l	Mg mg/l	B mg/l	Br mg/l	Li mg/l	TDS (g/l)	δD	δ ¹⁸ O	⁸⁷ Sr/ ⁸⁶ Sr
River Blanc	2005	RB2	70,3	6,79	595	153	74,6	156	360	46	38	4,5				1,17			
River Blanc	2005	RB3	92,8	8,31	2450	49	45,8	194	1331	119	72	<0,5	28,8	9,1	2,55	4,22	10,1	5,3	0,70465
River Blanc	2005	RB4	72,5	4,3	443	<0,5	203	211	273	32	36	6,9				0,93			
Roseau River	2005	RR1	84,5	7,3	1787	2005	25,3	184	1060	93	92	2	22,1	7,5	2,2	3,27	9,6	5	0,70467
Roseau River	2005	RR2	80,2	6,57	1182	305	29,5	221	745	59,9	58,6	10,3	14,4	4,2	1,72	2,4	-6,0	-1,0	
Trois Pitons R.	2005	3PR	50,1	6,3	679	198	18,2	126	455	33,6	35,3	6,2	8,1	2,5	0,98	1,4			
River Blanc	1985	DM15	96,6	7,97	2663	76,9	48	279	1579	202	52,9	0,48	9,7				-5,9	-0,1	
River Blanc	1985	DM29	95,3	7,95	2627	49,4	27,9	291	1549	176	66,9	0,24	20,4				-4,7	0,1	
Valley Desolation	2005	VD1	96,5	6,9	5892	32	93,8	291	868	208	1867	117	53,3	0,1	1,11		5,2	5,4	0,70456
Boiling Lake area	1985	DM17	85,6	3,6	1903	n.d.	68,2	279	851	178	260	21,6	27,5				6	3,6	
Boiling Lake area	1985	DM26	96,6	6,45	6308	36,6	96	284	1039	286	2441	178	40,7				-4,4	3,2	

Neutral sodium chloride waters (up to 93°C; 220 mg/l SiO₂; 4 g/l TDS) are discharged in Lower River Blanc and close to its junction with Roseau River and Trois Pitons River, i.e. at the lowest elevation in the field (Figure 3). They are mature waters and some of them plot on the full equilibrium line in the Na-K-Mg diagram of Giggenbach (1988). The Na-K geothermometer (Fournier, 1979) and the K-Mg geothermometer (Giggenbach, 1988) converge to a reservoir temperature of 200 -225°C (Figure 5). Fluid-rock interactions at high temperature are also supported by their isotopic composition with a shift in $\delta^{18}\text{O}$ from the local WML (Figure 6) and their strontium isotopic ratio ($^{87}\text{Sr}/^{86}\text{Sr}=0.7045\text{--}0.7049$) which is similar with andesitic rocks. In the Cl-Na, Cl-K, Cl-B, Cl-Li and Cl-Br binary diagrams (Figure 6, Figure 7, Figure 8), these sodium chloride waters belong to a same trend controlled by a mixing between a Cl-rich geothermal end-member and Cl-poor shallow groundwaters. In some diagrams, old analyses from BRGM (1985) like DM15 and DM 29 have not been considered because they deviate from the general trend. The chloride vs. magnesium binary plot shows a negative correlation, which is consistent with a progressive mixing and re-equilibration upon cooling causing an increase in magnesium content (Figure 7). According to this diagram, the chloride concentration of the Cl-rich geothermal end-member should be around 3000 mg/l. From the silica-chloride diagram (Figure 7), the silica concentration of this Cl-rich geothermal end member having a chloride concentration of 3000 mg/l should be similar to the silica content of sample DM29, i.e. about 291 mg/l. Based on the quartz geothermometer of Fournier (1977), the calculated quartz temperature of the Cl-rich end-member (190°C) is lower than the temperature range estimated previously from the Na-K-Mg diagram of Giggenbach (1988). It could be viewed as an evidence of cooling.

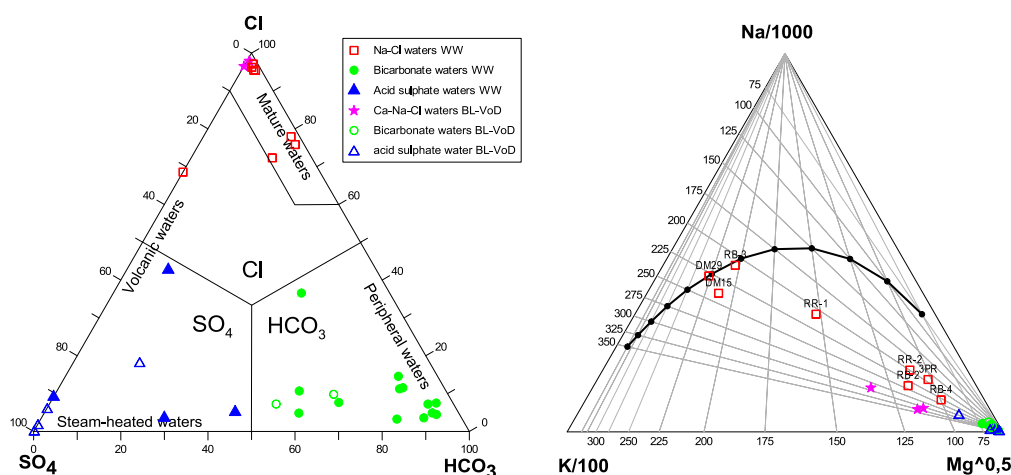


Figure 5 : Cl-SO₄-HCO₃ ternary diagram and Na-K-Mg ternary diagram of Giggenbach (1988) with full equilibrium line based on the Na-K geothermometer of Fournier (1979). WW: Wotten Waven and Laudat area; BL-VoD: Boiling Lake and Valley of Desolation area.

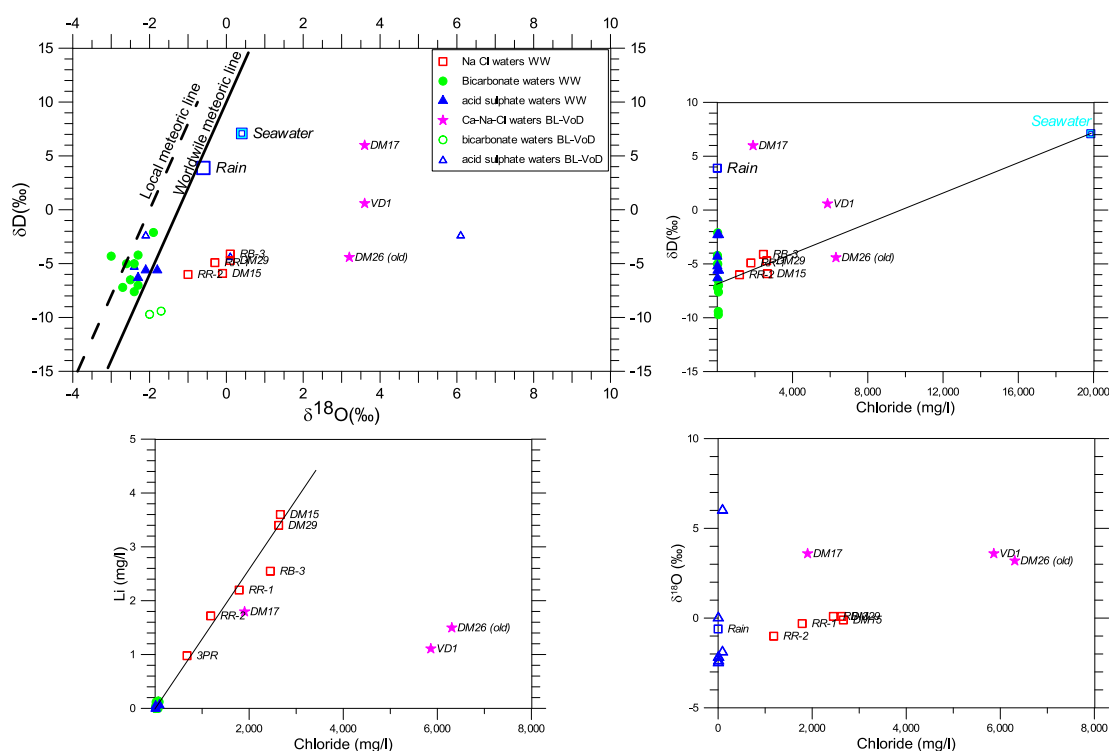
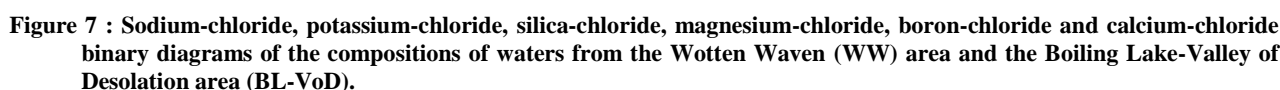


Figure 6 : δD - $\delta^{18}\text{O}$, δD -chloride, $\delta^{18}\text{O}$ -chloride and Lithium-chloride binary diagrams of the surface thermal waters of the Wotten-Laudat area (WW) and the Boiling Lake – Valley of Desolation area (BL-VoD).



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diagrams (Figure 8 and Figure 6) suggest a possible relationship between the Wotten Waven Na-Cl waters and these Ca-rich Na-Cl waters. The lack of equilibrium of the Ca-Na-Cl waters in the Na-K-Mg diagram (Figure 5) precludes the use of this graphic tool. The quartz geothermometer of Fournier (1977) gives 190°C while the Na-K geothermometer of Fournier (1979) gives 305°C. The lower temperature of the quartz geothermometer could be viewed as an evidence of cooling. The area of the Boiling Lake and Valley of Desolation deserves additional geochemical surveying in order to better characterize the waters and gases discharged by hot springs and steam vents and their relationships with Na-Cl waters discharged in Wotten Waven.

Bicarbonate waters are discharged as warm springs in River Camelia (Figure 3) and at greater distance from Wotten Waven (Trafalgar, Titou Gorge, Micotrin). They are representative of shallow aquifers heated by gas (CO₂) and steam above deep geothermal reservoir. The most striking feature is probably the occurrence of warm springs discharging bicarbonate waters on the flanks of the Micotrin lava dome at an elevation of 750-800 m.a.s.l., i.e. 500 meters higher than the main area of surface manifestations in River Blanc (Sulphur Springs). They could be indicative of a zone of fluid upflow from a deep geothermal reservoir. Warm and hot springs discharging acid-sulphate waters are found in the two main areas of surface manifestations in association with the solfataric activity (Sulphur Springs in Wotten Waven and Boiling Lake – Valley of Desolation).

4. GEOPHYSICAL MT SURVEY

A geophysical survey (MT, AMT, resistivity, gravity) had been done by BRGM in the 1980's (BRGM, 1984). In 2008, new MT and gravity surveys were done by BRGM (Baltassat et al., 2008). The purpose was to update the results obtained in the 1980's. It was also to evidence a connection at the reservoir depth between Wotten Waven and the Boiling Lake – Valley of Desolation, which are separated by 5 km (Figure 9). Because this region is covered with tropical forest and inaccessible to vehicles, it was difficult to survey the vicinity of the Valley of Desolation, and MT soundings were mainly restricted to the region of Wotten Waven and Laudat. Only a few have been sited along the trail to the Boiling Lake (Figure 9). Thirty-two MT soundings were measured over a surface of 25 km². Because of a low natural MT signal in addition to very low solar activity and strong cultural noise in the area of investigation, the signal-to-noise ratio was extremely low over the whole surveyed area in the low-frequency range. Although attempts were made to improve the signal-to-noise ratio by increasing the recording duration and filtering uncorrelated noise using remote references, most soundings show unusable data below 1 hertz. As a consequence, the penetration depth was limited to 400 m to 1000 m depending on the underground resistivity, and some soundings with poor data have to be removed from the interpretation. The results are presented on Figure 9 and Figure 10.

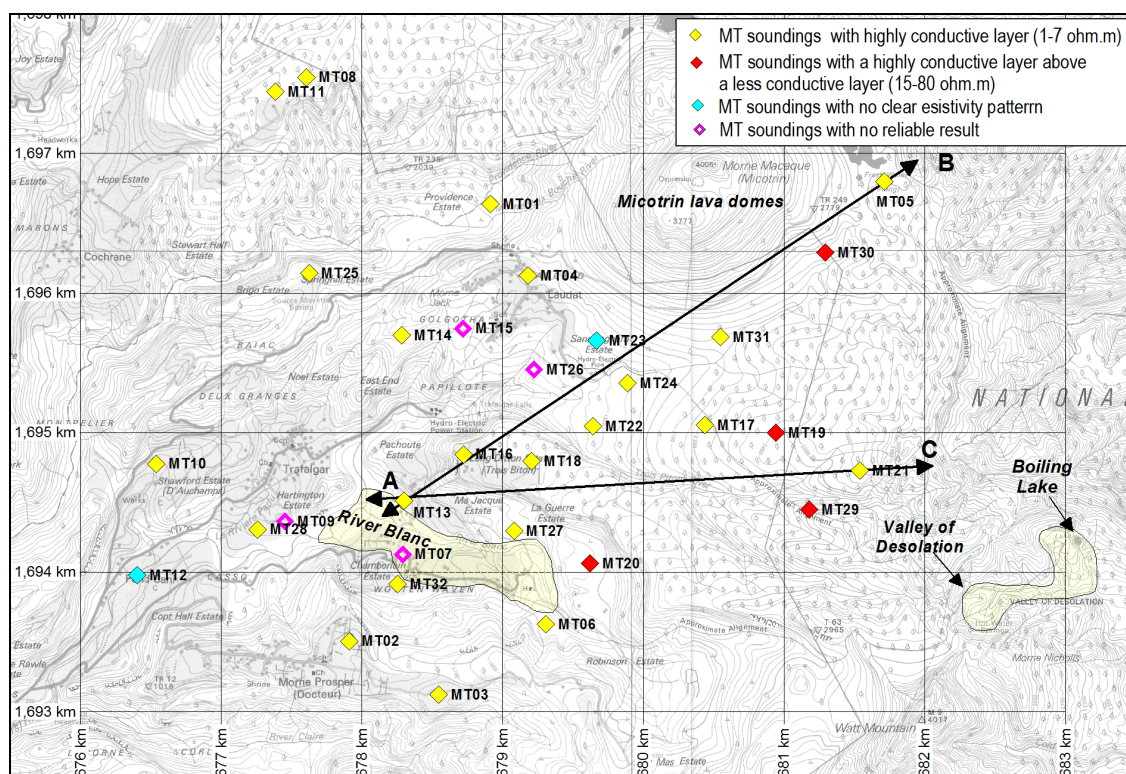


Figure 9 : Map showing the location of the 32 magnetotelluric (MT) soundings done by BRGM (Baltassat et al., 2008).

Most of the MT soundings recorded a highly conductive layer (less to 1 ohm.m to 7 ohm.m) as represented on Figure 9. This conductive layer is very close to the surface in the vicinity of the surface manifestations of River Blanc at Wotten Waven (Figure 10). Eastward in the area of Laudat, this conductive layer is overlain by 300 m to 450 m-thick, shallow resistive layer (20 to 700 ohm.m) which probably corresponds to the ignimbrite formations observed in the cliffs of the Trafalgar Falls. This highly conductive layer is ascribed to the development of an argillic cap-rock. Below it, a few soundings located in the central and eastern part of the surveyed area show a less conductive layer, with resistivity ranging from 15 to 80 ohm.m (Figure 9 and Figure 10). This range of resistivity is commonly observed in many high temperature geothermal reservoirs. On Figure 10, the thickness of the highly conductive layer appears to be about 250-350 m. This figure is also common for cap-rock in high temperature geothermal systems. Therefore, the MT survey provides a resistivity pattern in the central and eastern parts of the studied area, which is fully

consistent with the existence of a high-temperature geothermal system. At soundings MT19 and MT29, the bottom of conductive layer is at 500-600 m depth which corresponds to an elevation of 150-250 m.a.s.l.. In the areas of Laudat and Wotten Waven, the limited depth of penetration of the MT soundings unfortunately prevents to observe the existence of the less conductive layer below the highly conductive, argillic cap-rock. However, the abundance of surface manifestations is a good indicator of the existence of the reservoir at depth below the argillic cap-rock.

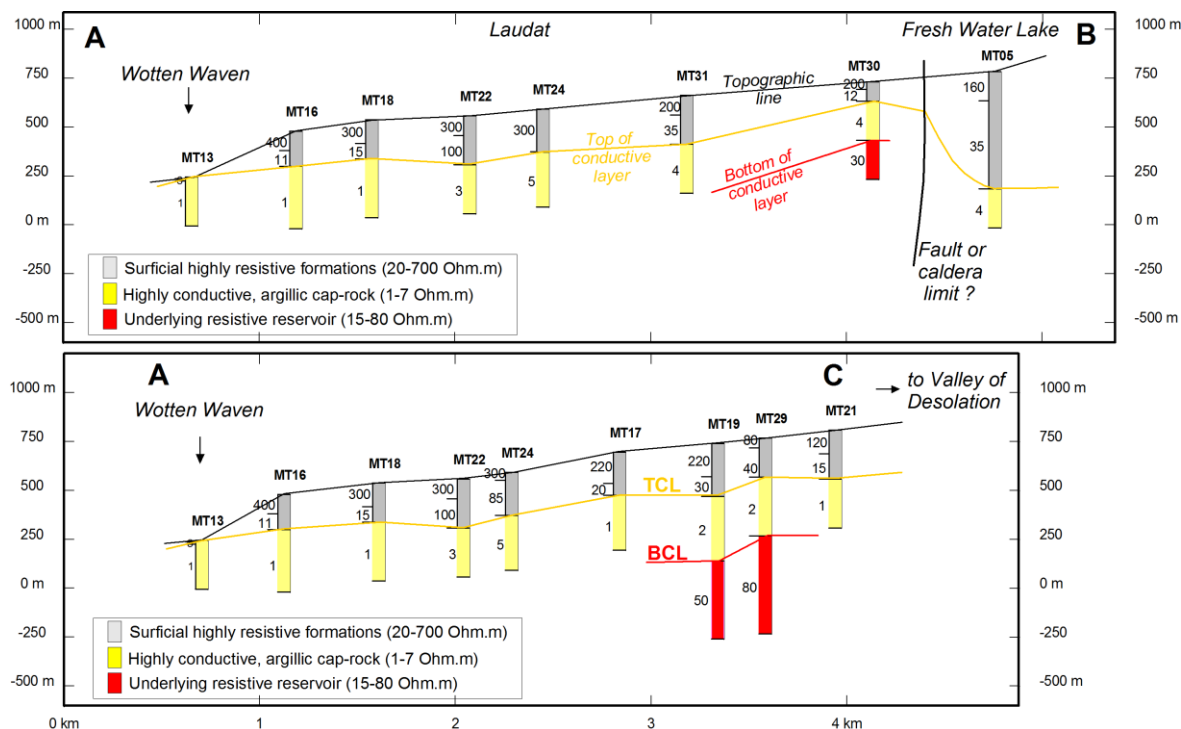


Figure 10 : Cross-sections along the A-B and A-C profiles depicted on Figure 9 showing the pattern of resistivity along MT soundings. TCL: top of conductive layer; BCL: bottom of conductive layer.

To the west and to the south of Wotten Waven, the limited depth of penetration also prevents from confirming the extension of the less conductive layer ascribed to the reservoir. The lack of surface manifestations as previously discussed and indicated on Figure 2 suggests that the reservoir is lacking. There is a geoelectric discontinuity between the MT soundings MT28, MT10 on one hand, the MT soundings MT02, MT32 and MT13 on the other (Figure 9), which suggests the existence of a N-S or NNW-SSE structure (fault, caldera border?). This structure may act as the permeability barrier quoted previously, in order to explain the lack of surface manifestation downstream of Wotten Waven. To the east, the results of the MT survey and especially the soundings MT19 and MT29 clearly suggest a connection between Wotten Waven and Boiling Lake – Valley of Desolation. There is another feature shown by the cross-section AB on Figure 10. The plunge of the Bottom of Conductive layer (BCL) between soundings MT31 and MT05 suggests the existence of a fault or a caldera limit.

As a conclusion, the MT survey fails to image the boundary of the high temperature reservoir as is often required, but it supports the model of a large reservoir extending between Micotrin lava dome, Wotten Waven, Boiling Lake and Valley of Desolation (Figure 9).

5. PROPOSED CONCEPTUAL MODEL

Lasne and Traineau (2005) proposed a first conceptual model of the Roseau Valley geothermal field. They assumed the existence of two distinct deep sodium chloride reservoirs (the Micotrin reservoir and the Boiling Lake reservoir) or a single large deep sodium chloride reservoir subject to mixing with seawater on its western margin. This initial model has been updated by Traineau et Lasne (2008) and Traineau et al. (2008), thanks to the additional data on fluid geochemistry and to the results of the MT survey. As mentioned previously, the latter supports the hypothesis of a large reservoir extending from Wotten Waven to Boiling Lake – Valley of Desolation. The Na-Cl waters and Ca-Na-Cl waters discharged in both areas might have a common origin, but would have evolved separately during their ascent to the surface. The higher contents of calcium and chloride in the Ca-Na-Cl waters discharged in the Boiling Lake –Valley of Desolation could be indicative of a strong degassing before they reach the surface.

The updated model which best fits with the results of the geological, geochemical and geophysical surveys is a deep sodium chloride reservoir with an upflow zone below the recent Micotrin lava domes, and lateral outflow zones to the southwest (Wotten Waven) and to the southeast (Boiling Lake – Valley of Desolation) as shown by Figure 11. The warm springs discharging bicarbonate waters on the flanks of the Micotrin lava domes (see Figure 2) well support the existence of an upflow zone. In this model, Wotten Waven is considered as the distal part of the southwest outflow zone and this is well consistent with the observed zoning of the surface manifestations (see Figure 3). The existence of a shallow steam-cap below the River Blanc area, in addition to the fluid ascent and its boiling is suggested in order to explain the abundance of surface manifestations. The existence of a shallow steam-cap below the Valley of Desolation is also suggested but not assessed.

The Boiling Lake – Valley of Desolation hydrothermal area is located high up at an elevation of 700-800 metres. The discharged Ca-Na-Cl waters appear to be less diluted and more representative of the deep Na-Cl parent fluid and it is consistent with a location closer to the upflow zone. The lateral fluid outflow zone probably develops along the inferred caldera border and crosses the central ridge thanks to a NNW-SSE fault (Figure 10).

The lack of surface manifestations in the Roseau valley west of Wotten Waven (except one occurrence of warm fluid and gas discharging in the Roseau River close to its junction with La Riviere Padu; Figure 2) suggest the existence of a hidden permeability barrier as reported on Figure 2, which could be either a fault zone or a caldera border. The extension of the reservoir to the south of the surveyed area (Robinson Estate, Du Mas Estate; Figure 9) is not dismissed but it is not confirmed by surface manifestations.

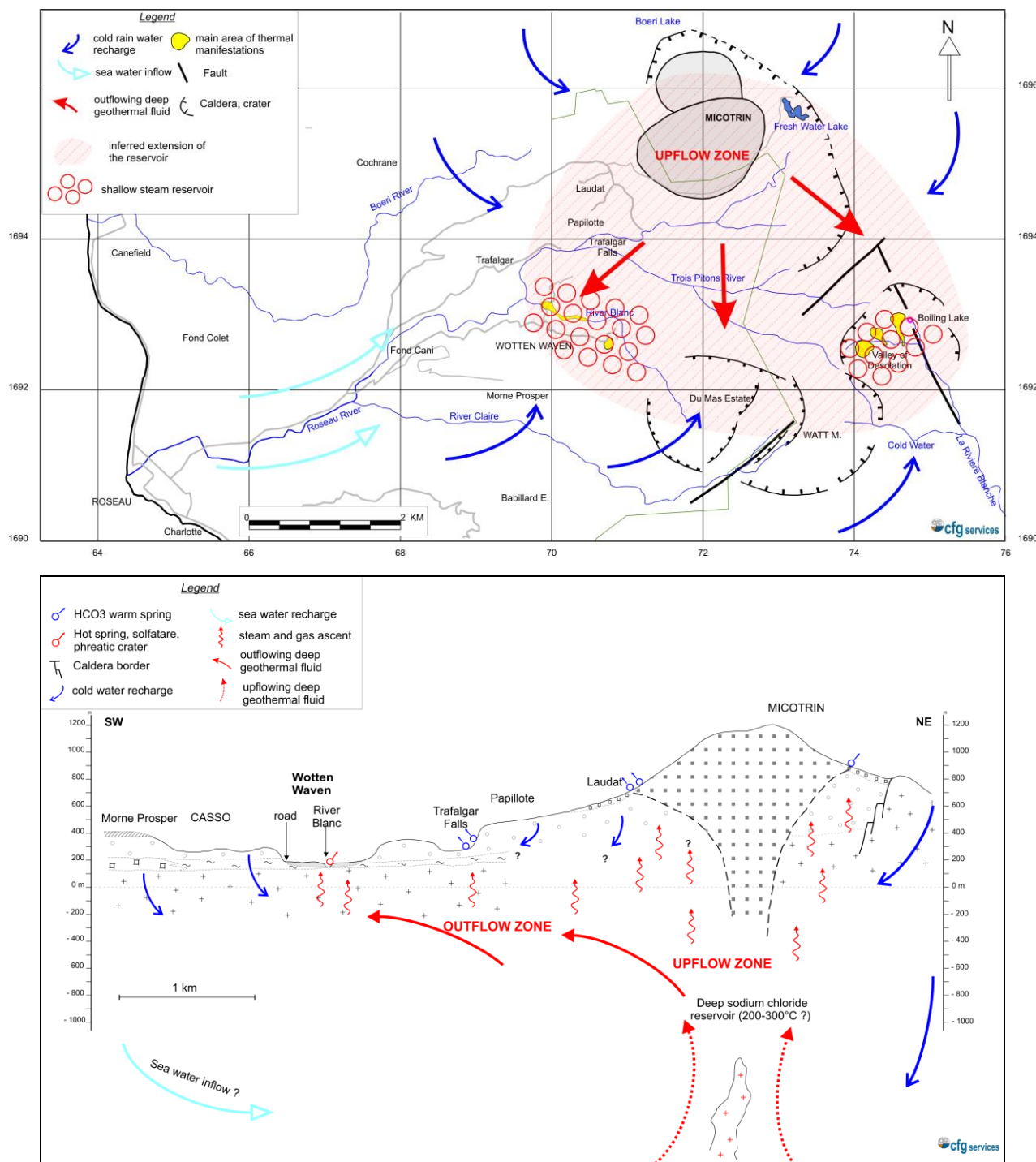


Figure 11 : Map and cross-section of the conceptual model of the Roseau valley geothermal field as proposed by Traineau and Lasne (2008).

6. RESOURCE ASSESSMENT

The estimated size of the reservoir is tentatively delineated on Figure 12. It corresponds to the area where high temperature fluids are thought to be encountered at depth and represents about 18 km². It is mainly based on the conceptual model. It includes the area

of the expected deep high temperature reservoir and upflow zone located below the Morne Micotrin lava dome. It includes also the outflow zones towards southwest (Wotten Waven) and southeast (Boiling Lake and the Valley of Desolation). The depth of the reservoir (as defined by the bottom of the conductive layer) has been estimated from the resistivity data provided by the MT survey (Baltassat et al., 2008). The top of the reservoir is thought to be very shallow in the River Blanc valley (300 to 500 m deep). To the northeast, it is increasing in parallel with the topography. In the vicinity of Laudat, the top of the reservoir is expected to be about 700-800 m deep. Around the Fresh Water Lake, it is assumed to be around 1000 m deep.

Temperature of the deep Na-Cl reservoir is not well constrained. According to the Na-K geothermometer of Fournier (1979) applied to the Ca-Na-Cl waters of the Valley of Desolation (sample VD-1), its maximum temperature could be around 300°C. Na-Cl waters discharged in Wotten Waven have been equilibrated at lower temperature (200-225°C).

Figure 12 proposes a preliminary design of the development of the Roseau Valley geothermal field. On a first basis, it is considered that the development should be focused on the western half of the reservoir which has been delimited. The eastern half appears to be located within the boundary of the National Park; it is a remote area without any existing road access. The main zone of interest extends south of the Micotrin lava dome on a large, gentle-slope plateau covered with forest. It is accessible through the narrow paved road linking Roseau to Laudat. It appears to be the most suitable area for siting production wells and large capacity power plants. Its eastern border corresponds to the boundary of the National Park. Directional wells can be drilled to exploit the deep high temperature fluids across the border if allowed by the Park regulations. Its southern border corresponds to the ignimbrite cliff overhanging the River Blanc valley. The Trafalgar Falls represent its western boundary. The road from Laudat to Fresh Water Lake lines it to the north. The elevation of this gentle-slope plateau ranges from 500 to 800 m.a.s.l. from west to east. At its southwestern margin, the right-hand side of the River Blanc valley (from Pachoute Estate to Robinson Estate; Figure 12) is considered as the most suitable reinjection zone for brines produced at higher elevation in the Laudat-Micotrin plateau. It could also be viewed as suitable for a small-size capacity power plant with a few production and reinjection wells. Its elevation ranges from 250 to 450 m.a.s.l. The Fond Cani-Casso area is considered to be at the margin of the high temperature reservoir. This area could be tentatively ascribed to fluid reinjection. However, the benefits in terms of pressure recovery could be limited.

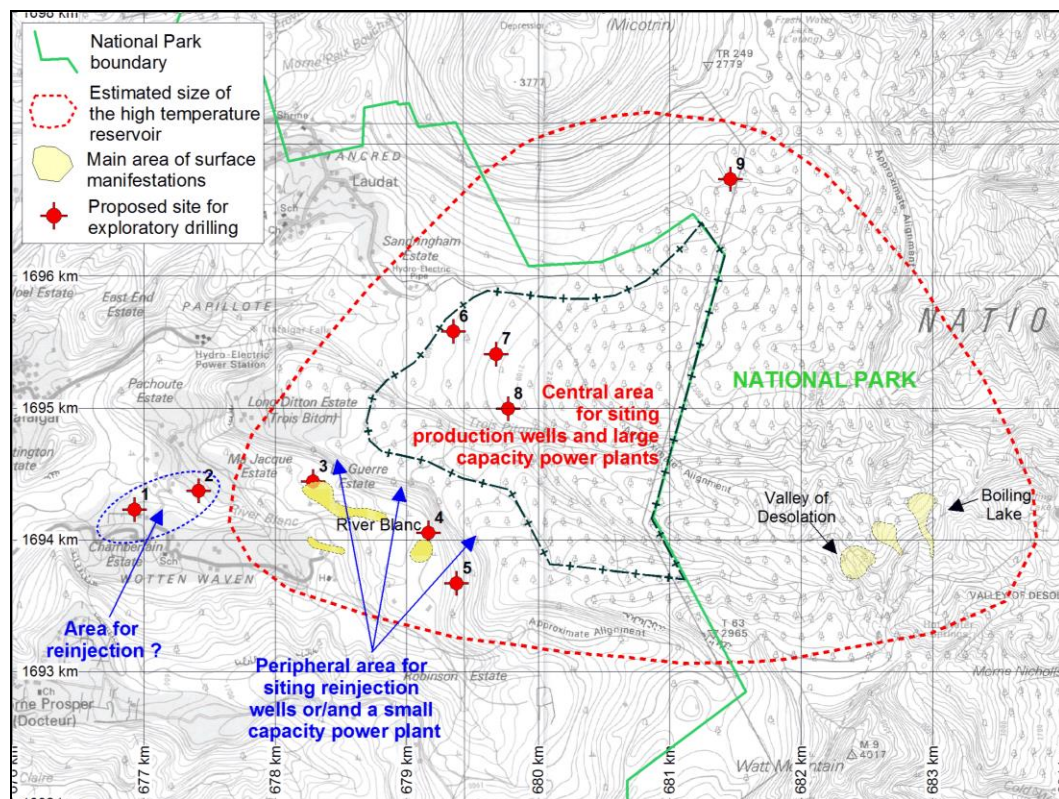


Figure 12: Estimated size of the Roseau valley geothermal reservoir and preliminary design proposed for the field development (from Traineau et al., 2008).

A preliminary assessment of the reservoir capacity has been conducted, by estimating the number of wells that could be sited within the western half of the reservoir. About 50 vertical and directional wells could be accommodated, if considering an average well spacing of 500 m on a 15 km² area. Assuming a ratio between production wells and reinjection/unsuccessful wells of 3:2, and given an equivalent power capacity of 4 MWe per production well, a preliminary figure of 100-120 MWe is considered for the reservoir capacity. It is in agreement with the rule of thumb of a power density of 10 MWe per km², which is observed for reservoirs with temperature around 250°C. At least sites for exploratory drilling have been proposed (Figure 12). They are scattered all over the reservoir area in order to check potential for production and reinjection.

Even if a conceptual model of the Roseau Valley high temperature geothermal system has been proposed, there are still some unknowns which have not been addressed by the successive surveys. The distribution and coverage of MT survey are incomplete.

Moreover, due to low natural MT signal, distribution of resistivity at depth remains unclear. An additional MT survey would be beneficial to cover larger area and to investigate deeper. The geochemical survey was also incomplete. There is no data on gas in fumaroles and solfataras. The structural analysis was not able to evidence main faults. As a consequence, it was not possible to define clear drilling targets for exploratory wells.

6. CONCLUSIONS

The Roseau Valley geothermal field can be considered one of the high potential geothermal prospects in the Caribbean. The abundance of surface manifestations and the chemistry of waters discharged are supporting the existence of a high temperature reservoir with temperature ranging between 200°C to 300°C. Despite its limited depth of penetration, the MT survey provided a clear resistivity pattern for some soundings which points out a highly conductive, argillic cap-rock and an underlying less conductive reservoir. However, there are some uncertainties which still exist after the last survey done in 2008. First, the area of the Boiling Lake and Valley of Desolation warrants an additional geochemical survey, in order to better characterize the waters and gases discharged by hot springs and steam vents, and their relationships with Na-Cl waters discharged in Wotten Waven. This area, together with the area north and east of the Micotrin lava domes, also deserves additional surveys (geological, geochemical, geophysical) in order to check the extension of the reservoir on the “Atlantic side”. This is a remote area covered with tropical forest and without vehicle access. According to the testimonies of local residents, there are no surface manifestations except the Boiling Lake and Valley of Desolation.

Another challenge will be to combine the development of the geothermal field with the Dominican environmental constraints and tourism. The Roseau valley is one of the most visited places in Dominica, thanks to the thermal manifestations of Sulphur Springs. Some small-scale facilities for spas have been developed here in the last ten years. The exploitation of the geothermal reservoir in the Wotten Waven area should take care of the flow rate of natural manifestations.

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