

## CO<sub>2</sub> Emissions from the Krafla Geothermal Area, Iceland

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### ABSTRACT

Geothermal resource utilization although widely accepted as a clean energy source, has also contributed to the decreasing of air quality due to hydrogen sulphide and carbon dioxide emissions. Several studies have shown that CO<sub>2</sub> emissions from geothermal/volcanic systems occur naturally and in some cases these natural emissions exceed the amount of CO<sub>2</sub> emitted from the geothermal power plant utilizing the geothermal resource. This study is carried out to quantify the natural CO<sub>2</sub> soil flux emissions from the Krafla geothermal field, see the relation between soil gas emission and the structural geology and compare the results to the CO<sub>2</sub> emissions from the Krafla power plant. The results of this study show that the total CO<sub>2</sub> flux from soil degassing is approximately 14.13 t d<sup>-1</sup> for a survey area of 2.5 km<sup>2</sup>, a positive correlation between CO<sub>2</sub> soil flux emissions and the structural geology of the area, and the CO<sub>2</sub> emission from natural sources exceeds the emission from the power plant by approximately 3 times.

### 1. INTRODUCTION

#### 1.1 Background

Geothermal resource utilization in Iceland has shown significant benefits in the form of supplying clean, renewable energy and has made the country's capital Reykjavik one of the cleanest cities in the world. Although widely accepted as a clean energy source, geothermal utilization especially high temperature utilization for generating electricity has contributed to decreasing air quality due to hydrogen sulphide and carbon dioxide emissions despite the fact that these are much lower than emissions from fossil fuel combustion power plants (Giroud and Arnórsson, 2005). The latter is known as a greenhouse gas and with the implementation of the Kyoto protocol and awareness of global warming becoming stronger amongst environmentalists and the general public on a global scale, this issue has become more important. Several studies have shown that CO<sub>2</sub> emissions from geothermal/volcanic systems occur naturally and in some cases these natural emissions exceed the amount of CO<sub>2</sub> emitted from the geothermal power plant utilizing the geothermal resource (e.g., Seaward and Kerrick, 1996; Delgado et al., 1998; Bertani and Thain, 2002). A study in the Lardarello field in Italy has shown a noticeable and measurable decrease in the natural release of CO<sub>2</sub> from the ground as a result of geothermal power development (Bertani and Thain, 2002), while a study in New Zealand has shown that the exploitation of the Wairakei system significantly increased diffuse surface heat flow, which if heat flow is considered as a proxy for CO<sub>2</sub> emissions it could lead to conclusions that the exploitation has increased the natural CO<sub>2</sub> emissions (Sheppard and Mrocze, 2004). In Iceland, a study in the Reykjanes geothermal area in SE-Iceland has

concluded that the planned power plant (which is now operating) will significantly increase the CO<sub>2</sub> emissions from the geothermal system (Fridriksson et al., 2006). In addition they found that the natural emissions were predominantly soil diffuse emissions (Table 1) as had been suggested by other workers in other areas (e.g. Favara et al. 2001, Sorey et al. 1998, Evans et al. 2002, Gerlach et al. 2001). Studies of CO<sub>2</sub> emissions from geothermal power plants and natural geothermal activity in Iceland have also been conducted by Ármannsson et al. (2005), as shown in Table 2.

**Table 1. Reykjanes. Different Conduits (modified from Fridriksson et al., 2006).**

	CO <sub>2</sub> (t d <sup>-1</sup> )	Steam (t d <sup>-1</sup> )	Heat flow (MW)
<b>Soil</b>	13.5	4150	108
<b>Steam vents</b>	0.23	72	1.9
<b>Steam heated pools</b>	0.15	46	1.2

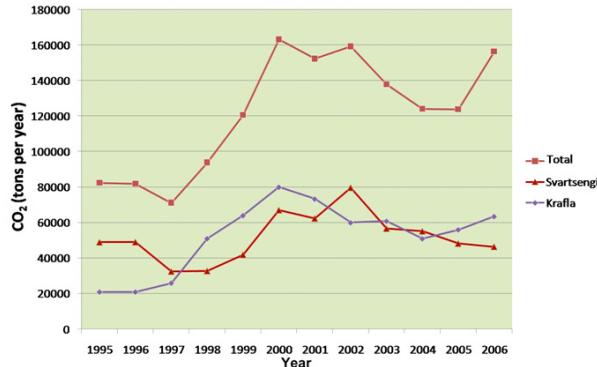
CO<sub>2</sub>: Soil: 97.3%; Steam vents: 1.6%; Pools: 1.1%

**Table 2. CO<sub>2</sub> and S (expressed as SO<sub>2</sub>) emissions per kWh from Iceland's major geothermal power plants in 2000 (Ármannsson et al., 2005).**

Plant	From electricity generation only		From electricity and heat production	
	CO <sub>2</sub> (g kWh <sup>-1</sup> )	S as SO <sub>2</sub> (g kWh <sup>-1</sup> )	CO <sub>2</sub> (g kWh <sup>-1</sup> )	S as SO <sub>2</sub> (g kWh <sup>-1</sup> )
Krafla	152	23	152	23
Svartsengi	181	6	74	2
Nesjavellir	26	21	10	8

So summing up these studies there is still some controversy on this issue, and we could say that different areas will show different behaviour in CO<sub>2</sub> emissions affected by geothermal utilisation and they should be assessed for each area. The main controversy is whether the emissions from geothermal plants is an addition of gas to the atmosphere or whether it is just a transfer from natural emissions to plant emissions. There is some evidence that in vapour dominated systems the emissions are large to start with but then will decrease and if averaged over some years they can be treated as just a transfer while the same does not apply to

liquid dominated systems. Ármannsson and Fridriksson (2008) presented results for  $\text{CO}_2$  emissions from two geothermal plants, Krafla and Svartsengi, along with the total emissions from all geothermal plants in Iceland (Figure 1).



**Figure 1:  $\text{CO}_2$  from Icelandic geothermal plants 1995-2006 (Ármannsson and Fridriksson 2008)**

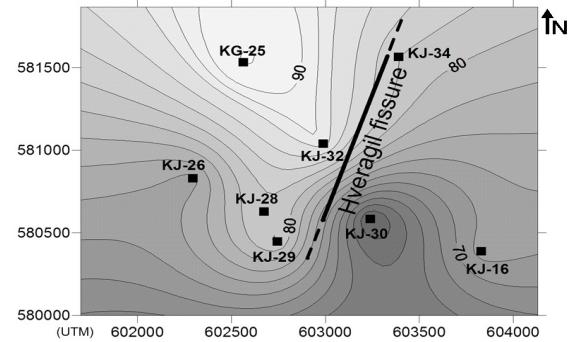
In the early 1990s a steam pillow had developed and was produced from in the Svartsengi area and  $\text{CO}_2$  emissions had increased substantially. In 1996 however the gas emissions were waning again in spite of continued production. In 1998-2000 more wells were drilled into the steam pillow and a substantial increase in gas emissions was observed. By 2003 the emissions had decreased again. In Krafla several wells were drilled in 1997-1999 to increase the installed power of the power plant from 30 to 60 MW. Most of the wells were high enthalpy or close to being vapour dominated. This resulted in a considerable gas increase but this was decreasing in 2003 although production had not been decreased. This supports the view that the gas content of high enthalpy steam will decrease after an initial increase. The increase observed for the total emissions from geothermal plants in Iceland in 2006 is due to the commissioning of two new geothermal power plants, Hellisheiði and Reykjanes.

Ármannsson et al. (2007) studied the concentration of carbonate in cuttings from the drilled part of the Krafla area and obtained an apparently inverse relationship between the amount of carbonate fixed in rock and soil diffuse  $\text{CO}_2$  emissions, also suggesting that a substantial part of the  $\text{CO}_2$  flux from the magma is bound in the rocks close to the surface (Figure 2).

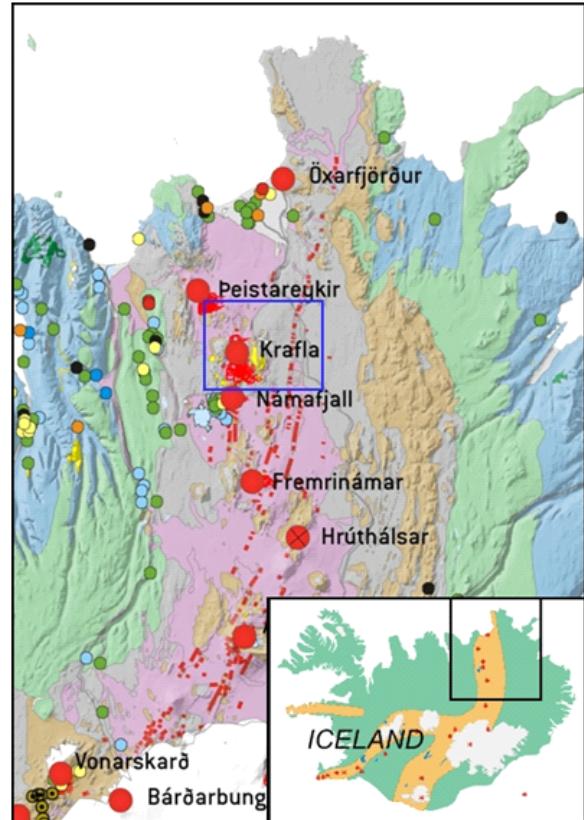
## 1.2. $\text{CO}_2$ Emissions from the Krafla Geothermal Area

The Krafla geothermal area is located within the neovolcanic zone in NE- Iceland (Figure 3). It consists of the Krafla central volcano and a 100 km long North-South transecting fissure swarm. It has a 10 km wide caldera that was formed around 100,000 years ago by a violent ryholitic tuff-forming eruption. Krafla has been the source of many rifting and eruptive events during the Holocene, including two in historical time. The Mývatneldar eruptions (the “Mývatn fires”) in 1724 began with a great volcanic explosion which formed the crater Stóra - Víti. In the following years, a series of earthquakes and eruptions occurred in the vicinity of Krafla mountain. The greatest eruption took place in 1729, when lava flowed from Leirhnjúkur mountain down to Mývatn lake. Eldhraun, the lava field formed during the eruptions. This system was last active between 1975 and 1984 when lava erupted from, and to the north of the central volcano, and dykes were injected along most of the fissure zone (Saemundsson,

1991; Björnsson et al., 1979). This episode is known as the Krafla fires. This event has increased the gas emissions (dominantly  $\text{CO}_2$ ) from the area significantly due to a magmatic intrusion.



**Figure 2: Krafla  $\text{CO}_2$  fixed in bedrock. (contours tons/m<sup>2</sup>) (Ármannsson et al. 2007)**

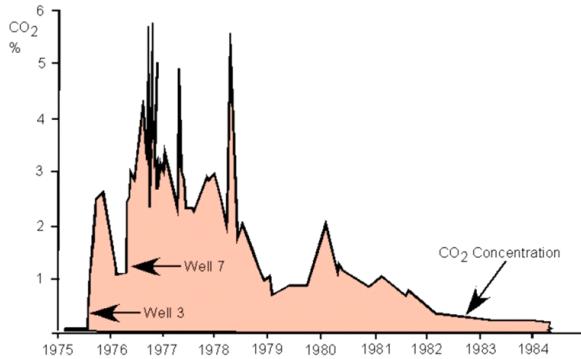


**Figure 3: Map of the Krafla geothermal area**

Drilling started in Krafla in 1974 but the power plant was commissioned in 1978 then only producing 7 MW. Drilling was halted due to the Krafla fires but was resumed in 1980 until 1982 after which the plant produced 30 MW (Ármannsson et al. 1987). The second turbine was commissioned in 1999 after which the plant has produced (Júlíusson et al. 2005).

Some wells in the Leirbotnar field which in the beginning of the Krafla fires was the only field that has been drilled were blocked due to the formation of deposits of pyrite and pyrrhotite in the course of the magmatic gas passage to the surface. The gas concentration has been carefully monitored and the pattern has been similar for wells in the affected Leirbotnar field, i.e. a maximum in 1977/1978, a

secondary maximum in 1980 and a steady decline since (Figure 4). In Figure 4 changes that were first observed in well 3 were subsequent to that well's collapse followed in the nearby well 7, both wells being in the Leirbotnar field (Ármannsson et al., 1989).



**Figure 4: Changes in CO<sub>2</sub> concentrations in wells 3 and later the neighbouring well 7 during the Krafla fires 1975-1984**

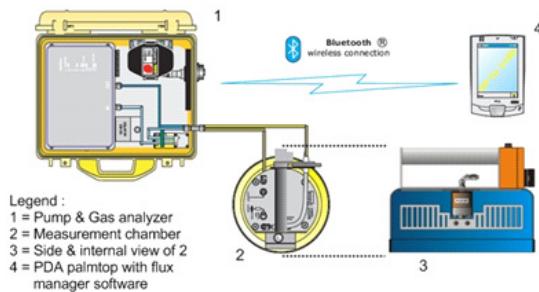
In a previous study of CO<sub>2</sub> budget of the Krafla geothermal system carried out by Ármannsson et al. (2007) soil CO<sub>2</sub> flux emissions and CO<sub>2</sub> concentration in drill cuttings were determined. The result of their study shows that the mean flux of the geothermal population is about 115 g m<sup>-2</sup>d<sup>-1</sup> and it emanates from about 10% of the total area. The total CO<sub>2</sub> flux from the eastern Krafla caldera is about 120 kt/yr and about 70% of that is of geothermal origin (Ármannsson et al. 2007).

In this study the research area is north and west of where Ármannsson et al. (2007) had already collected data is extended by an additional area of approximately 1 km<sup>2</sup>. In the area of study only CO<sub>2</sub> soil flux is taken into account since there have been no boreholes present there. It is hoped that the results of this study will complement the previous study and show the relation between soil gas emissions and the structural geology of the Krafla geothermal area.

## 2. METHODOLOGY

### 2.1. CO<sub>2</sub> Flux Measurements

The CO<sub>2</sub>-flux measurements are carried out directly with a closed-chamber CO<sub>2</sub> flux meter from West Systems. The flux meter is equipped with a LICOR LI-820 single-path, dual wavelength, non-dispersive infrared gas analyser (Figure 5). The flux meter has a 3.06.10<sup>-3</sup> m<sup>3</sup> internal volume. The flux measurement is based on the rate of CO<sub>2</sub> increase in the chamber, the measurement lasts for approximately 2 minutes at each location.



**Figure 5: Flux meter from West systems (model)**

CO<sub>2</sub> flux through soil was measured over a rectangular grid with intervals of 25 m N-S and 50 m E-W with some exclusion in areas not suitable for measurement (Figure 6). Data from previous measurements were also included in the analysis adding a total area of 2.5 km<sup>2</sup> and 3095 measurement points (Figure 6).

Figure 7 shows typical results of a CO<sub>2</sub>-soil flux measurement. Initially the CO<sub>2</sub> concentration inside the cell is constant at about 700 to 900 ppm but after approximately 40 seconds the CO<sub>2</sub> concentration starts to increase linearly with time. The slope of the curve defined by the CO<sub>2</sub> concentration as a function of time is a measure of the CO<sub>2</sub> flux through the soil. Other parameters that need to be accounted for when evaluating the flux from the concentration as a function of time are temperature inside the chamber, air pressure and the internal volume of the system.

The relationship between these parameters and the flux is defined by the following equation:

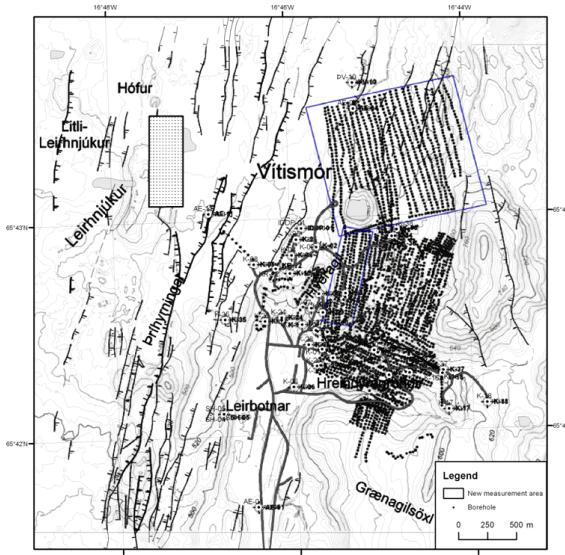
$$K = \frac{86400 \cdot P \cdot V}{10^6 \cdot R \cdot T_k \cdot A} \quad (1)$$

Where   
 $K$  = Accumulation chamber factor  
 $P$  = Barometric pressure [ mBar (HPa) ]  
 $R$  = Gas constant [ 0.08314510 bar LK<sup>-1</sup> mol<sup>-1</sup> ]  
 $T_k$  = Air temperature [ Kelvin ]  
 $V$  = Chamber net volume [ m<sup>3</sup> ]  
 $A$  = Chamber inlet net area [ m<sup>2</sup> ]

The dimensions of  $K$  is :

$$K = \frac{\text{moles} \cdot \text{meter}^{-2} \cdot \text{day}^{-1}}{\text{ppm} \cdot \text{sec}^{-1}}$$

The values of  $K$  can be obtained from the table provided by the equipment manufacturer.



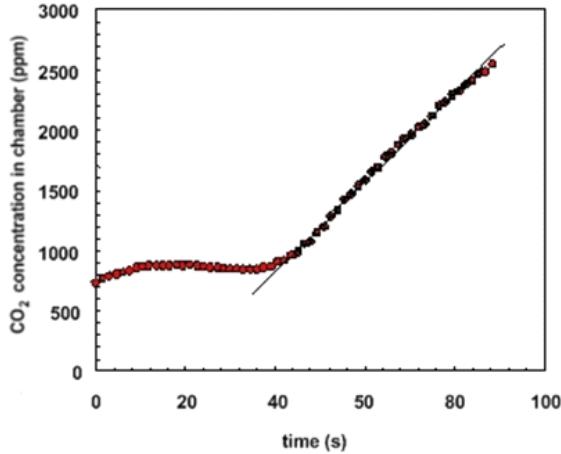
**Figure 6: CO<sub>2</sub> flux measurement points in the Krafla geothermal area**

### 2.2. Sampling Procedure

The flux meter chamber is pressed firmly against the ground and loose soil is packed (if necessary) around the outside. This is done to seal the measurement unit and

prevent atmospheric air from entering the system and affecting the measurement.

The CO<sub>2</sub> flux measurements are only conducted in dry weather conditions that have prevailed preferably for 2 days. This is to avoid potential effects of water saturation of the soil pores.



**Figure 7: Typical raw data from a soil flux measurement**

The appropriate distance between measurement points varies but the general rule of thumb is that at least three or four measurements are needed in order to define the anomalies. So if the widths of the anomalies are of the order of 100 m the grid spacing can be of the order of 25 to 30 m between points. Flux measurements on a grid allow the construction of diffuse soil degassing maps. These maps can also be used to discover “hidden” geothermal systems for which hydrothermal surface features (e.g., hot springs, elevated ground temperatures, hydrothermal alteration) are not present (Lewicki and Oldenburg, 2004).

### 2.3. Data Interpretation

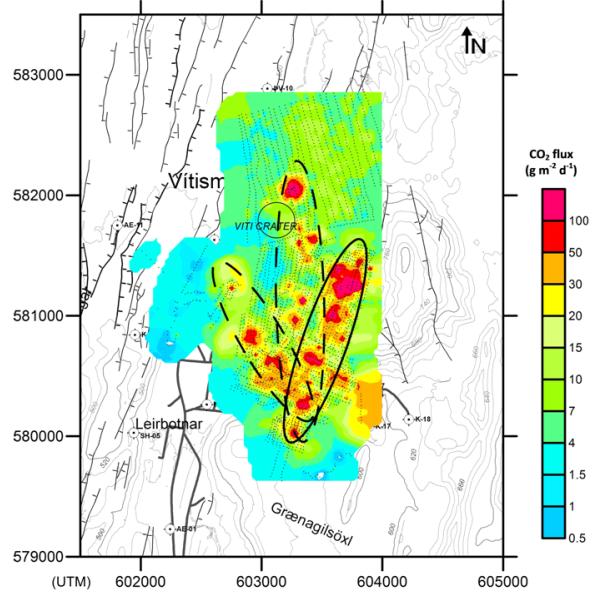
Collected CO<sub>2</sub> flux data is in ppm/sec, and is converted into  $\text{g m}^{-2}\text{d}^{-1}$  using equation 1. The data is then analyzed using the graphical statistical analysis (GSA) method of Sinclair (1974) to identify different populations within the samples and distinguish between background and anomalous CO<sub>2</sub> flux populations, determine the mean flux value and the standard deviation of the population (Ármannsson et al. 2007). The procedure of Sinclair is based on a detailed analysis of the distributions in probability plots.

## 3. RESULTS AND DISCUSSION

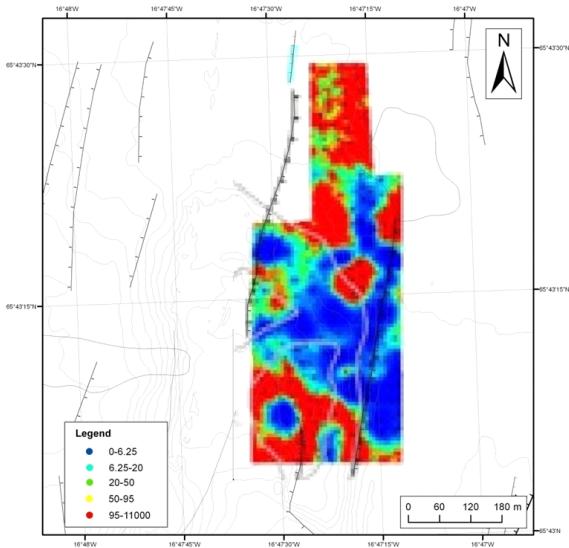
The CO<sub>2</sub> flux measurements were carried out according to the procedures described above. The CO<sub>2</sub> flux contours are shown in Figure 8 below. Previous results for the Leirhnúkur area are shown in Figure 9.

From the collected data, the GSA method of Sinclair (1974) was used to partition the population. This method has been successfully applied to the results of CO<sub>2</sub> flux campaigns in order to both separate background populations from anomalous CO<sub>2</sub> flux populations (i.e., where the fluxes originate in deep volcanic-hydrothermal CO<sub>2</sub>) and to compute the total CO<sub>2</sub> output, and relative uncertainties, from the different sources active in surveyed areas (Fridriksson, et al., 2006).

The logarithmic probability plot in the Krafla geothermal area (Figure 8) shows that the entire data set has a polymodal density distribution. The plot forms a curve with two inflection points (marked with arrows) dividing the populations into three theoretical populations with log normal distributions, A, B, and C with proportions ( $f_i$ ) of 25%, 71.5%, and 3.5% respectively.

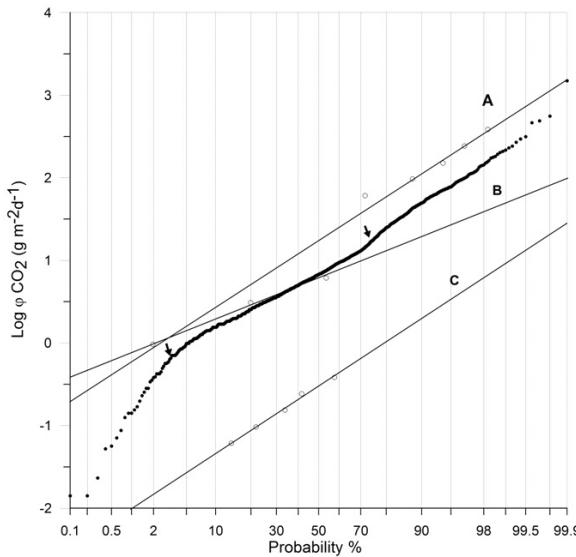


**Figure 8: Contour map of CO<sub>2</sub> flux from measured points**



**Figure 9: Results of soil diffuse CO<sub>2</sub> flux measurements in the Leirhnúkur area (after Ármannsson et al. 2007)**

This result is then used to determine the mean ( $M_i$ ) and standard deviation ( $\sigma_i$ ) of each population by plotting the 50% cumulative probability to the log  $\varphi$  CO<sub>2</sub> intersecting each population line for the mean, and the 84% subtracted by 50% plotted values for the standard deviation. The results are reported in Table 3.



**Figure 10: Logarithmic probability plot of CO<sub>2</sub> through soil at Krafla. Arrowhead shows inflection points dividing the population into three at 3.5% and 75% cumulative probability. A theoretical population of A, B, and C obtained following the graphical statistical method of Sinclair (1978)**

Because the means and standard deviations refer to the logarithm of the values, the estimation of the mean soil  $\varphi$  CO<sub>2</sub> value ( $MN_i$ ) and the 95% confidence interval of the mean for each population is found by means of the Sichel's *t*-estimator (David, 1977). The area covered by each population ( $S_i$ ) is estimated by multiplying  $f_i$  with the total surveyed area ( $S = 2,500,000 \text{ m}^2$ ). The CO<sub>2</sub> output from each population is then obtained by multiplying  $S_i$  with  $MN_i$ . Finally we calculate the total CO<sub>2</sub> output from the surveyed area by summing the CO<sub>2</sub> output from each population. It can be seen from Table 2 that the total CO<sub>2</sub> output is 26.33 t d<sup>-1</sup> (9610 t yr<sup>-1</sup>) with the estimated maximum and minimum values 27.5 and 24.9 t d<sup>-1</sup> respectively. From this output, about 54% is from population A (geothermal origin) and 46% is from background emissions.

These results are lower than those reported by Ármannsson et al., (2007) for which the calculated mean flux from the geothermal population about 115 g m<sup>-2</sup> d<sup>-1</sup> and the total CO<sub>2</sub> flux from the eastern Krafla caldera about 120 kt yr<sup>-1</sup>. This is not surprising since the new surveyed area generally showed a low CO<sub>2</sub> flux except for some points North-East of the Viti lake crater where there are surface manifestations of geothermal activity such as steam vents and mud pools. Background concentrations in the range of

95% confidence interval of the estimations are given in parenthesis 0.5-15 g m<sup>-2</sup> d<sup>-1</sup> (Figure 5) were mostly observed in the other areas. The sampling points in their study are also skewed towards areas with visible geothermal manifestations while in this study the sampling points are more uniform covering a certain area not taking into account whether there are geothermal manifestations or not. The pattern of the CO<sub>2</sub> flux shows a NE-SW trend which confirms the fault trending of the Krafla geothermal area (parallel with the Hveragil fissure).

The estimated output from the surveyed area is only  $0,0096 \cdot 10^9 \text{ kg yr}^{-1}$  which is relatively low compared with most of the other volcanic and geothermal areas shown in Table 4.

**Table 4. CO<sub>2</sub> output from some volcanic and geothermal areas (Source: Ármannsson et al., 2005, original sources can be found there).**

Area	CO <sub>2</sub> output ( $10^9 \text{ kg yr}^{-1}$ )
Pantelleria Island, Italy	0.39
Vulcano, Italy	0.13
Solfatara, Italy	0.048
Ustica Island, Italy	0.26
Popocatepetl, Mexico	14.5-36.5
Yellowstone	10-22 <sup>a</sup>
Mammoth Mountain, USA	0.055-0.2
White Island, New Zealand	0.95
Mt. Erebus, Antarctica	0.66
Taupo Volcanic Zone, New Zealand	0.44
Furnas, Azores, Portugal	0.01
Mid-Ocean Volcanic System	30-1000
Total	200-1000

<sup>a</sup> Diffuse degassing only

**Table 3. Estimated Parameters of the Partitioned Populations and Diffuse CO<sub>2</sub> Output.**

Population	$f_i$ (%)	$M_i \pm \sigma_i$	$MN_i (\text{g m}^{-2} \text{d}^{-1})$	No. of points	$S_i (\text{m}^2)$	Source	$FCO_2 (\text{t d}^{-1})$
A	25	$1.27 \pm 0.62$	22.6 (23.9-21.4)	774	625,000	Hydrothermal	14.13 (14.94-13.38)
B	71.5	$0.8 \pm 0.39$	6.81 (7.03-6.45)	2213	1,787,500	Background	12.17 (12.57-11.53)
C	3.5	$-0.5 \pm 0.65$	0.39 (0.44-0.35)	108	87,500	Background	0.0341 (0.0385-0.0306)
Total	100			3095	2,500,000		26.33 (27.5-24.9)

If we take this result and do an extrapolation to the estimated total area of natural CO<sub>2</sub> degassing in Krafla which is about 50 km<sup>2</sup> (Ármannsson, 2003) this will give a result of 192,210 t yr<sup>-1</sup>. The CO<sub>2</sub> emissions from the Krafla power plant were reported as of 63,500 t in the year 2006 (Ármannsson and Fridriksson, 2008). We can see that the amount of the natural emissions exceeds the emissions from the power plant by approximately 3 times. These natural amount only encompass emissions from soil and not from focussed degassing and other natural conduits which will add a small amount to this natural emission if we assume that soil diffuse emissions are the dominant natural source of CO<sub>2</sub> emissions in the area.

## CONCLUSIONS

The soil CO<sub>2</sub> flux concentration contour is conforming the fault trends of the area (NE-SW) and high concentrations are found in areas where surface manifestations are present.

The study shows that CO<sub>2</sub> emissions in geothermal areas occur naturally even without visible surface manifestations through soil diffuse degassing and in this particular case the amount well exceeds the CO<sub>2</sub> emission from the power plant utilizing the geothermal energy in the area. The amount of soil diffuse CO<sub>2</sub> flux from geothermal origin is estimated around 14.13 t d<sup>-1</sup> for a survey area of 2.5 km<sup>2</sup>, while the total emissions from natural sources are estimated around 192.21 kt yr<sup>-1</sup> for a 50 km<sup>2</sup> area compared to the CO<sub>2</sub> emission from the Krafla power plant about 63 kt yr<sup>-1</sup>.

It is beneficial to see the CO<sub>2</sub> emissions trend with time from both natural sources and geothermal utilization to see if there is a relation between them, a periodical monitoring of CO<sub>2</sub> soil flux emissions would give us an opportunity to understand better the impact of geothermal utilization on CO<sub>2</sub> emissions. This could be accomplished by placing automatic continuous monitoring stations which measures soil CO<sub>2</sub> flux and various environmental parameters which can potentially affect the soil gas flux at selected sites, along with a more detailed periodic measurements at fixed points that is repeated several times per year. Granieri et al., (2002) reported the results of a continuous CO<sub>2</sub> soil flux measurement in the Solfatara crater (Phlegraean Fields, Italy) for a period of 4 years (1998-2002) through a combination of an automatic continuously operating station at a selected site and periodic measurements of flux over an array of sites.

## REFERENCES

Ármannsson, H., 2003: CO<sub>2</sub> emission from geothermal power plants. Proceedings of the International Geothermal Conference IGC-2003 "Multiple Integrated Uses of Geothermal Resources" Reykjavík, S12, 56-62.

Ármannsson, H., and Fridriksson, Th., 2008: Geothermal energy and greenhouse gases: natural emissions due to production, and fixation of carbon dioxide in bedrock (in Icelandic). ÍSOR, Reykjavík, webpage: [www.isor.is/pdf/ársfundarindi%202008 %20ÍSOR %20HÁ%20PF.pdf](http://www.isor.is/pdf/ársfundarindi%202008 %20ÍSOR %20HÁ%20PF.pdf).

Ármannsson, H., Benjamínsson, J., and Jeffrey, A.W.A., 1989: Gas changes in the Krafla geothermal system, Iceland. Chemical Geology, 76, 175-196.

Ármannsson, H., Fridriksson, Th., Gíslason, S.R., 2008: Carbon dioxide emissions from Icelandic geothermal systems and the possibility of capturing carbon dioxide by dissolution and reaction with basalt, ISOR, presented at Guanajuato, Mexico.

Ármannsson, H., Fridriksson, Th., Kristjánsson, B.R., 2005: CO<sub>2</sub> emissions from geothermal power plants and natural geothermal activity in Iceland. Geothermics, 34, 286-296.

Ármannsson, H., Fridriksson, Th., Wiese, F., Hernández, P., and Pérez N., 2007: CO<sub>2</sub> budget of the Krafla geothermal system, NE-Iceland. Proceedings of the 12th International Symposium on Water-Rock Interaction 2007, Taylor & Francis Group, London, 189-192.

Ármannsson, H., Gudmundsson, Á., and Steingrímsson, B.S., 1987: Exploration and development of the Krafla geothermal area. Jökull, 37, 12-29.

Bertani, R. and Thain, I., 2002: Geothermal power generating plant CO<sub>2</sub> emission survey. IGA News, 49, 1-3.

Björnsson, A., Johnsen, G., Sigurdsson, S., Thorbergsson, G., and Tryggvason, E., 1979: Rifting of the plate boundary in North Iceland 1975-1978. J. Geophys. Res., 84, 3029-3038.

David, M., 1977: Geostatistical ore reserve estimations. Developments in Geomathematics 2, Elsevier, Amsterdam, 364 pp.

Delgado, H., Piedad-Sánchez, N., Galván, L., Julio, P., Alvarez, J.M., Cárdenas, L., 1998: CO<sub>2</sub> flux measurements at Popocatépetl volcano: Magnitude of emissions and significance (abstract). EOS Trans. Am. Geophys. Union, 79 (45), 926.

Evans, W.C., Sorey, M.L., Cook, A.C., Kennedy, B.M., Shuster, D.L., Colvard, E.M., White, L.D. and Huebner, M.A. 2002: Tracing and quantifying magmatic carbon discharge in cold groundwaters: lessons learned from Mammoth Mountain, USA. J. Volc. Geoth. Res., 114, 291-312.

Favara, R., Giannanco, S., Inguaggiato, S. and Pecoraino, G. 2001: Preliminary estimate of CO<sub>2</sub> output from Pantelleria Island volcano (Sicily, Italy): evidence of active mantle degassing. Applied Geochemistry, 16, 883-894.

Fridriksson, T., Kristjánsson, B.R., Ármannsson, H., Margrétardóttir, E., Ólafsdóttir, S., and Chiodini, G., 2006: CO<sub>2</sub> emissions and heat flow through soil, fumaroles, and steam heated mud pools at the Reykjanes geothermal area, SW Iceland. Applied Geochemistry, 21, 1551-1569.

Gerlach, T.M., Doukas, M.P., McGee, K.A., and Kessler, R., 2001: Soil efflux and total emission rates of magmatic CO<sub>2</sub> at the Horseshoe Lake tree kill, Mammoth Mountain, California, 1995-1999. Chem. Geol., 177, 101-116.

Giroud, N., and Arnórsson, S., 2005: Estimation of long-term CO<sub>2</sub> and H<sub>2</sub>S release during operation of Geothermal Power Plants. Proceedings of the World Geothermal Congress 2005, Antalya, Turkey, CD, 6 pp.

Granieri, D., Chiodini, G., Marzocchi, W. and Avino, R. 2003: Continuous monitoring of CO<sub>2</sub> soil diffuse degassing at Phlegraean fields (Italy): influence of environmental and volcanic parameters. Earth and Planetary Science Letters, 212, 167-179.

Júlfusson, B.M., Pálsson, B. and Gunnarsson, Á., 2005: Krafla power plant in Iceland – 27 years of operation. Proceedings of the World Geothermal Congress 2005, Antalya, Turkey, 4 pp.

Lewicki, J. L., and Oldenburg, C.M., 2004: Strategies for detecting hidden geothermal systems by near-surface gas monitoring. Lawrence Berkeley National Laboratory, University of California, webpage: [repositories.cdlib.org/lbnl/LBNL-56895/](http://repositories.cdlib.org/lbnl/LBNL-56895/).

Saemundsson, K., 1991: The geology of the Krafla system. In: Gardarsson, A., and Einarsson, Á. (eds.), The natural history of Lake Mývatn (in Icelandic). The Icelandic Society of Natural Sciences, Reykjavík, 25-95.

Seaward, T.M. and Kerrick, D.M., 1996: Hydrothermal CO<sub>2</sub> emission from the Taupo Volcanic Zone, New Zealand. *Earth Planet. Sci. Lett.* 139, 105-113.

Sheppard, D. and Mroczek, E., 2004: Greenhouse gas emissions from the exploitation geothermal systems. *IGA News*, 55, 11-13.

Sinclair, A.J., 1974: Selection of threshold values in geochemical data using probability graphs. *J. Geochem. Explor.* 3, 129–149.

Sorey, M.L., Evans, W.C., Kennedy, B.M., Farrar, C.D., Hainsworth, L.J. and Hausback, B. 1998: Carbon dioxide and helium emissions from a reservoir of magmatic gas beneath Mammoth Mountain, California. *J. Geophys. Res.*, 103, 15,303-15,323.