RESOURCE EVALUATION AND DEVELOPMENT PLANS FOR A 120 MW GEOTHERMAL POWER PLANT ON MILOS ISLAND, GREECE

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ABSTRACT

Five deep wells have been drilled on the Island of Mixos, Greece, identifying a high-temperature, high-enthalpy goothermal reservoir. The thermodynamic properties of the fluid, and the porosity and thickness of the formation suggest a fluid and heat storage capacity that could support a 60 MWe power plant for 85 years or a 120 MWe for half that time.

The existing five wells can deliver 180 t/h of steam at 10 bar ab's pressure, capable of generating a maximum electric power output of slightly less than 20 MWe.

This paper describes the **geology**, the drilling and the well testing results pertaining to the five wells, and discusses the reservoir potential for a 60 MWe geothermal power plant.

INTRODUCTION

Milos Island, a member of the Cyclades Group in the Aegean Sea, is part of the active Aegean Volcanic Arc of which the islands of Santorini and Nisyros are also members. Economides et al. have described the geothermal field on Iisyros in an associated paper in this volume. Fig. 1 from that paper shows the location of Milos Island within the volcanic arc.

Two exploratory wells, MA-1 and MZ-1, drilled in 1975-76 proved productive. The former was drilled in the Adamas area, while the latter was drilled in the Zephyria area. Three

other wells drifted during 1980-82 (M- \bar{i} , M-2 and M-3) all proved productive. Figure 1 positions the wells on the map of the island.

GEOLOGY

The Istand of Mitos may be represented by the generalized mode's shown on Figure 2 offered by Vrouei'. The schematic represents a Y-E cross-section of the island at the Patitude of the Zephyria plane. The mode2 shows an infittration zone in the west (associated with several Pava domes), a partial deep discharge towards the east, the development of a series of convective circuits within the reservoir (metamorphic complex), some upward leakages through the main fault systems, a cover made up of widespread hydrothermal alterations of a polygenic formation, and some local intrusions connected with the few domes outcropping in the south-eastern sector of the island. The ratter intrusions, however, represent a local hydrogeological disturbance to the deep aratem.

Because of the reasonably flat surface terrain and the proximity to population centers, the Eastern part of the Island, and especially the Zephyria-Agrilies and the Adamas regions were sited for the contemplated geothermal development.

DRILLING AID SUBSURFACE LITHOLOGY

Table 1 presents a summary of the well diameters and depths for the three newer wells (M-1, M-2, and M-3).

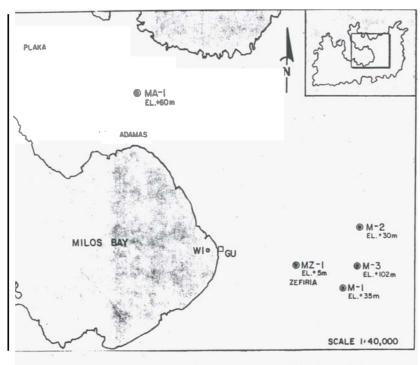


Figure 1. Location of Wells M1, M2, M3, MA-1, and MZ-1 on Milos

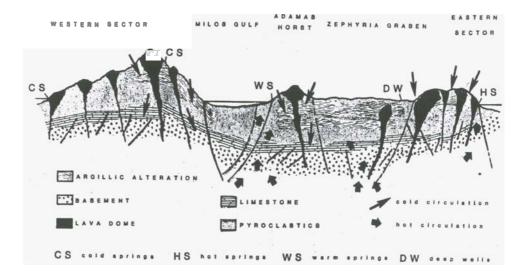


Figure 2. Sketch of Geothermal Model of the Island of Milos (From Reference 2).

TABLE 1 DRILLING STATISTICS FOR WELLS M-1, M-2, and M-3

Depth Well Diameter	M-1	M-2.	M-3
	1.180 m	1381	1017 m
24" 17 1/2" 12 1/4" 8 1/2"	0-99.5 m 99.5-399 m 399-903 m 903-1180 m	0-169 m 169-600 m 600-1381 m	0-186 m 186-637 637-1017 m

The sequence of terrains crossed by each well is summarized below.

M-1: A shallow interval of alluvia (0-20 m) is made of stratified flood deposits. From 20-60 flood deposits. From 20-60 m there is the chaotic formation made of merange the chaotic formation made of metange comprised of schists, micaschists, quartz, limestones and extremely altered materials. Beneath this terrain there is a thick metamorphic layer (60 m to T.D.) consisting of schist primarity of the green schist facies. Epidote is a prominent mineral in this complex.

 $\underline{\text{M-2}}$: The **alluvia** terrain is again a shallow terrain (0-15 m) composed of stratified flood desposits. A thin Eager (15-25 m) of grey tuff and other new volcanics is present. From 25-145 m, there is the chaotic formation consisting of a merange of various lithotypes. Metamorphic fragments (quartz, mica-schists, chroro-schists) are followed by volcanic fragments, some sedimentary brocks and hydrothermal material. Below 149 m (to T.D.) a thick metamorphic formation similar to the one under M-1 is found.

A volcanic formation (20-125 m) is characterized by intense hydrothermal alteration rendering much Of the original structure of the rock unrecognizable.

This layer is underlain by a sedimentary formation (125-173 m) formed by the Neogenic marine ingression. The formation contains carbonaceous quartz arenitea, sandstones, arenaceous carbonate rocks, arenit**ea**, limestones and dolomites and some shales.

Finally, the thick metamorphic layer observed in M-1 and W-2 is found below 173 m (to T.D.).

The terrain sequences found in the three wells is summarized in schematic form in Figure 3. A geoffolio history of the structures may then be reconstructed.

A pull-apart fault probably caused the block of rock benenth well $M\!-\!3$ to subside; followed by the marine ingression which led to the deposition of the Neogenic formation (125-173 m). Generally, the sediments are well sorted. Sometime fater in the geologic history, the magma from a moster rock body under the metamorphic rock complex intruded upward pushing the whole block upward depositing it above the neogenic formation by uplifting the overlying metange deposita. Another possibility is that magma formed a volcanic cone around the vent above the neogenic sediments. Then melange and affiliviar formations were deposited later above the cone the cone

The focation of the $M\!\!-\!\!3$ on the top of the hill forming the uplifted horst that borders the Zephyria plain suggests the presence Of two faults (see cross-section) around M-3. This hypothesis is further strengthened by the presence of neogenic sediments in M-3 and their absence in M-1 and M-2.

Considering the mineralogy and tithotogy of the rocks pierced by the wells, it can be suggested that the metamorphic trend is from M-2 to M-3 to M-1.

Presence of calcite mineral above 900 m and its absence at lower depths. can be directly related to the pattern of underground water movement.

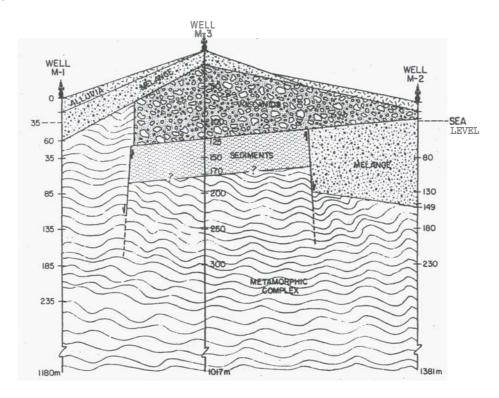


Figure 3. Terrain Sequences penetrated by wells M-1, M-2, and M-3. N-S Cross section.

FLOW AND TRANSIEST PRESSURE WELL-TESTING

The three new wells have been ffortested and Table 2 contains the pertinent flowrate and enthalpy data.

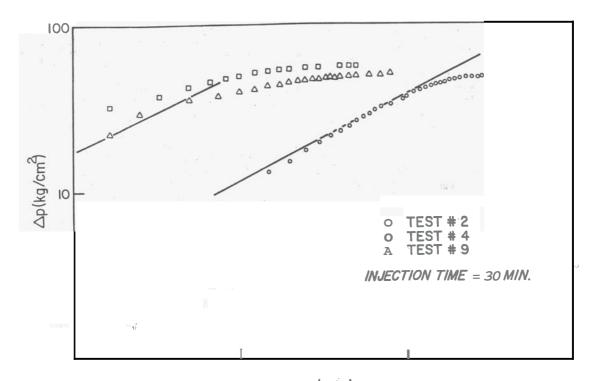
A Targe number of injection teets have beet done on att three wells. Cold seawater was itjected and the etsuits buildup of pressure ras attowed to fall off. Nine such teste were done for well M-1. The injection time for each war approximately 3Q minutes. Nine similar tests were done for H-2 and three more were done for H-3.

The tests were performed at various well intervals for the evafuetion of the recervoir transmissivity. Log-Yog graphs of these teats were utilized as diagnostic too's, features of Yog-Yog graphe are used extensively by reservoir engineers to identify reservoir characteristics (such as the presence of fractures) or

wellbore conditions. Figure 4 is a log-log graph of three injection fall-off tests from M-1, while Fig. 5 is a similar graph of three injection fall-off teets from M-3.

In general, moat of the data demonstrate the presence of fractures, as identified by the half-slope as identified by the harr-stops behavior on the Rog-kog graph evident at early time. The data trend then a frequent indicator of flattens out, a frequent indicator of double porosity systems. A "rule of thumb' known as the "double Ap rule" destities the commencement of the flattens out, thumb' known as the down in the straight Bine behavior in a semi-Togarithmic graph of the data.

According to the rule, data following two times the Ap from the deviation from the half stope will fall on a semi-log straight line. Unfortunately, most of the testa were not rut for e sufficiently long time for the appearance of the semi-Pog straight line as is evident from Figs. 4 and 5. Further, the ehort duration of the



At (min)
Figure 4.' Injection Fall-off Pressure Transients Tests for M-1

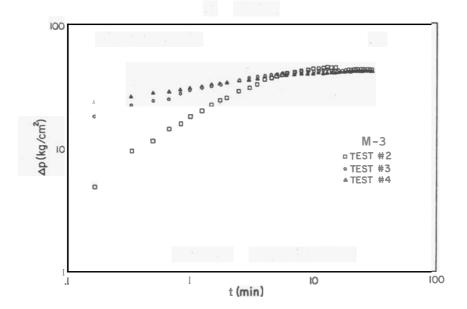


Figure 5. Injection Fall-off Pressure Transients Tests for M-3

TABLE 2		
ELL FLOW DATA		
<u>K-1</u> 119 t/h	H-2 47 t/h	$\frac{\text{M}-3}{126} \text{ t/h}$
1460 kJ/kg	2200 kJ/kg	1600 kJ/kg
0.35	0.71	0.42
41.7 t/h	33.4 t/h	52.9 t/h
4.6 MWe	3.7 MWe	5.8 MWe
900-T.D.	940-T.D.	900-T.D.
120,000 PPH	140,000 PPM	130,000 PPB
323°C	282°C ·	300°C+
\sim 120 kg/cm ²	$\sim 120 \text{ kg/cm}^2$	$\sim 120 \text{ kg/cm}^2$
	ELL FLOW DATA M-1 119 t/h 1460 kJ/kg 0.35 41.7 t/h 4.6 MWe 900-T.D. 120,000 PPH 323°C	ELL FLOW DATA M-1 119 t/h 1460 kJ/kg 0.35 0.71 41.7 t/h 33.4 t/h 4.6 MWe 3.7 MWe 900-T.D. 120,000 PPH 140,000 PPM

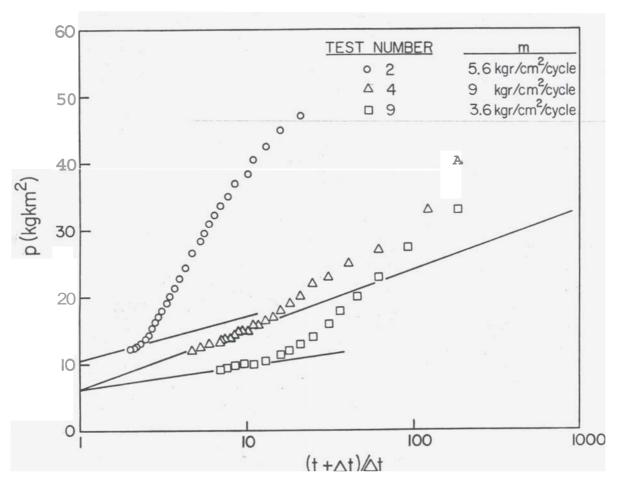


Figure 6. Semi-Logarithmic Analysis of Pressure Transients Data for Uell M-1

tests limits the value of the data f4r interpretation of the double porosity behavior.

The catculated storage capacity of well M-1 is 43 a (8 1/2" by 1180 m). A half hour injection at 100 m /h (as it was done for most tests) would barely unload the original wellbore contents into the formation.

Semi-Pog graphs of the data are presented in Fig. 6. Injection tests #4 and #9 were selected for analysis since they were done over roughly the same completion interval. The extrapolation of the time function (t+\Delta t)/\Delta to unity (infinite fa\%-off time, \Delta t) would result in the initial reservoir pressure. Hence, since the tests were done for the same depth, the extrapolation of these straight lines must converge at approximately the same value of pressure. The slopes for the two straight lines were thus obtained and were calculated as 9 kgr/cm2/cycle (for test #4) and 3.6 kgr/cm2/cycle (for test #9).

The pertinent equation for carcutating the permeability is:

$$k = \frac{526 \text{ q}\mu}{\text{mh}} \tag{1}$$

where k in mdarcys, q in $m^3/\text{hr},~\mu$ in cp and h in m.

The permeability values obtained from these two tests were $11.5\,$ md and $14\,$ md , respectively.

The wellbore drainage factor or skin effect may be calculated using the well known equation.

$$s = 1.151 \frac{p_{1hr} - p_{i}}{m} - \log \frac{k}{\phi \mu c_{t} r_{w}^{2}} + 3.11$$
(2)

where (kgr/cn^2) -1 and r_y in m.

Using Eq. 2, the skin effect fbr both tests was calculated as a = -7. The negative value for the skin further supports the notion that the next penetrates a fracture.

Injection test #2, done at a shafflower interval was also analyzed

providing a permeability value of 3.7 md and a skin factor of -6.

The Injection tests performed during the dritting operation provided a number of useful results about the formation and the wells. The diagnostic tog-log phots gave evidence of the presence of natural fractures, a widespread feature of gebthermal reservoirs elsewhere. The flattening out of the data trends gave indication of two porosity systems also evident in other geothermal formations. Analysis of three of these teste provided estimates for the formation permeability and the open hole wellocre condition. The negative value of the skin effect provided additional evidence of the penetration of natural fractures by the wells.

Short-time injection testa, utilizing alien to the formation fluids (seawater) at much lower tem erature would create other phenomena (such as thermal fracturing) that would mask or alter certain formation characteristics. At the present time, drawdown/buildup and interference well tests are contemplated in order to evaluate formation characteristics as well as the principal axes of permeability. The latter are extremely important in any future re-injection schemes which must be planned in conjunction with the development of the power plant.

RESERVE ESTIMATION

Assuming a maximum of 25 km² area% extent with a 1000 m thick reservoir (much of it via% be accessible through vertical fractures touching the well, not through drimming) a buck votume of 2.5 x 10 m² ray be calculated. Using a porosity value of 5%, the pore volume may then be calculated as 1.25 x 10 m². The figure specific volume at reservoir conditions (p=120 kg/cm²) is 1.53 x 10 m²/kg. Hence, the mass present is 1.25 x 10 m²/kg. Hence, the mass present is 1.25 x 10 m²/kg. Since 20,000 kg/hr or 9100 kg/hr have been traditionally used per 1 MWe, the total flow need for a 60 mWe, assuming 50% quantity is 1.09 x 10 kg/hr. Hence, the expected longevity of the reservoir is calculated as 85 years.

DEVELOPMENT PLANS

At the time of writing this paper, the Public Power Corporation is in the first phase of the development of the geothermal! reservoirs of Milos, and is in the process of contracting the design and installation of a small 1:5-2 MW condensing power plant by aid 1985. This pilot facility will provide a means for testing prototype separators, pipelines, turbines, condensers and re-injection schemes for problems associated with scaling and corrosion. In addition, extended fluid production and re-injection will provide more definitive data on the reservoir and on the sustained deliverability and injectivity characteristics of the wells.

The second phase will include the drilling of five new geothermal wells, testing and measurements, the study of a plausible gebthermal model and the determination of the total power capacity (expected to be about 60 MWe) resulting in the design, procurement and installation of a 60 MWe power plant by late 1909. The installation of the first network of submarine cables for the transmission of the generated power will also take place during this phase.

The third phase will include the drifting of approximately 20 new geothermal velts, appropriate testing and measurements aiming towards the procurement and installation of a

second 60 MWe power pfant by fate 1992. This phase will include the installation of a second network of submarine cables for the transmission 6f the additional 60 MWe.

Since the power output is expected to greatly outpace the demand for Milos or that of the islands in its immediate proximity, the installation of the submarine cables for power transmission to the Greek mainland is considered of utmost importance and is contemplated at the present time by the PPC.

REFERENCES

- 1. Economides, M.J., EntigEconomides, C.A., Koutroupis, N.,
 Spetiotis, G., Ungemach, P., and
 Yrousi, F.: "Pretiminary Assessment
 of the Geology and Reservoir
 Characteristics of the Geothermat
 Field on Misyros Istand, Greece",
 Proc. New Zealand Geothermat Workshop,
 1983.
- 2. Vrousi, F: Research and Development of Geothermal Recoursee in Greece: Present Status and Future Prospects, Public Power Corporation Report, Athens, 1983.

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