

TEMPERATURE GRADIENT MAP OF THE HAKKODA GEOTHERMAL AREA

OKUBO, Y. and MURAOKA, H.,

Geol. Surv. of Japan, Higashi 1-1-3, Tsukuba 305 Japan

Ninety-nine boreholes temperature log data in the Hakkoda geothermal area were gathered. The data were obtained from many kinds of boreholes of geothermal and mineral exploration and hot spring prospecting. Standing time of temperature logging of geothermal exploration was 24 hours or more, but most of hot spring data has no description of standing time. In addition, Depths of logging range from several tens to 1500 meters depth. Therefore, reliability of the data is not uniform. We divided the data into six classes based on the depth and condition of logging such as standing time and stability of the data. Considering reliability of the data, we drawn the temperature gradient contour lines by hand (Figure 1).

Background temperature gradient of the map ranges from $40^{\circ}\text{C}/\text{km}$ to $50^{\circ}\text{C}/\text{km}$, and high temperature gradient regions over $100^{\circ}\text{C}/\text{km}$ occur over the Hakkoda volcanoes (which is a Quaternary volcano group) and the Okiura and Aoni hot springs. The value of the background corresponds to the average temperature gradient of the backarc side of the Japanese Islands (Okubo et al., in press).

Figure 2 shows the relation between anion indexes and temperature gradients. High anion index values indicate high probability of heat sources. High temperature gradient zones roughly correspond to anion index value highs. Figure 3 shows the relation between gravity map and temperature gradients. Gravity highs occur around the Hakkoda volcanoes, while gravity lows occur to the west of the volcanoes. That is, a north-south trending gravity changing area extends on the west edge of the volcanoes. This suggests that faults or tectonic lines can occur. A high temperature zone extends north on the north-west of the volcanoes, which correlates well to the gravity changing area. Movement of geothermal fluid may be controlled by the faults or tectonic lines.

The map of temperature gradient certainly reveals the subsurface temperature between the ground surface and several kilometers depth and is appropriate to assess geothermal resources and to choose promising area in a reconnaissance stage. Borehole data of hot springs without any description of temperature logs are useful only if measured depths are sufficiently deep.

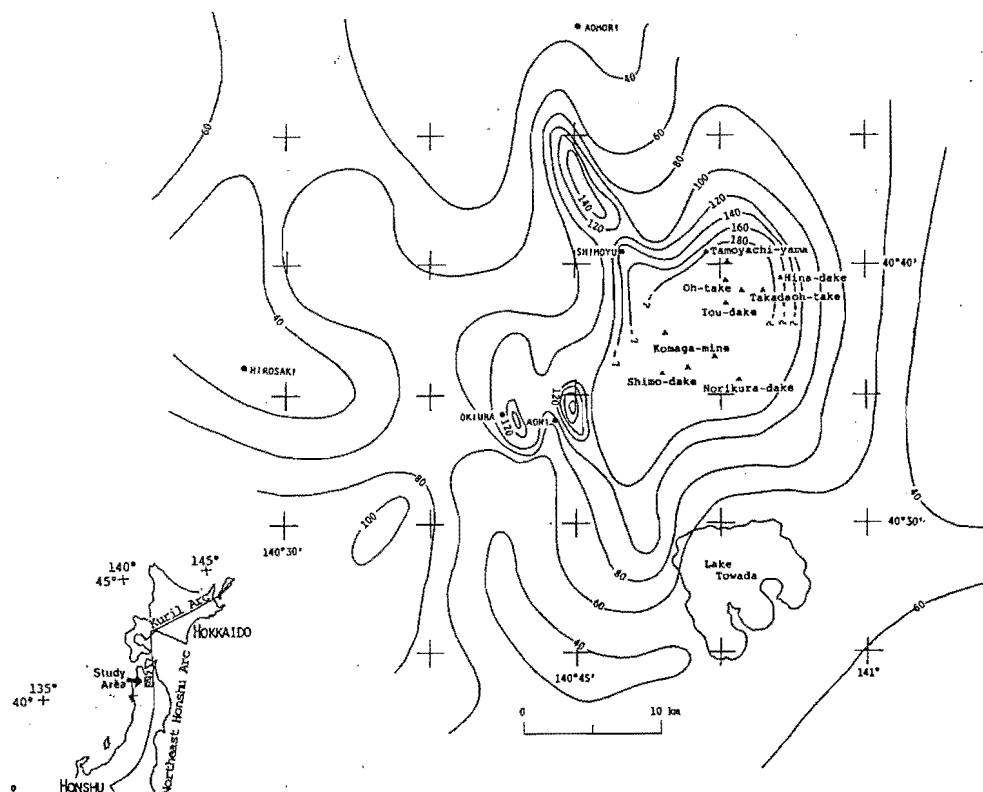


Fig.1 Temperature gradient map of the Hakkoda volcanoes. Contour interval is $20^{\circ}\text{C}/\text{km}$. Solid triangles indicate summits of the Hakkoda volcanoes.

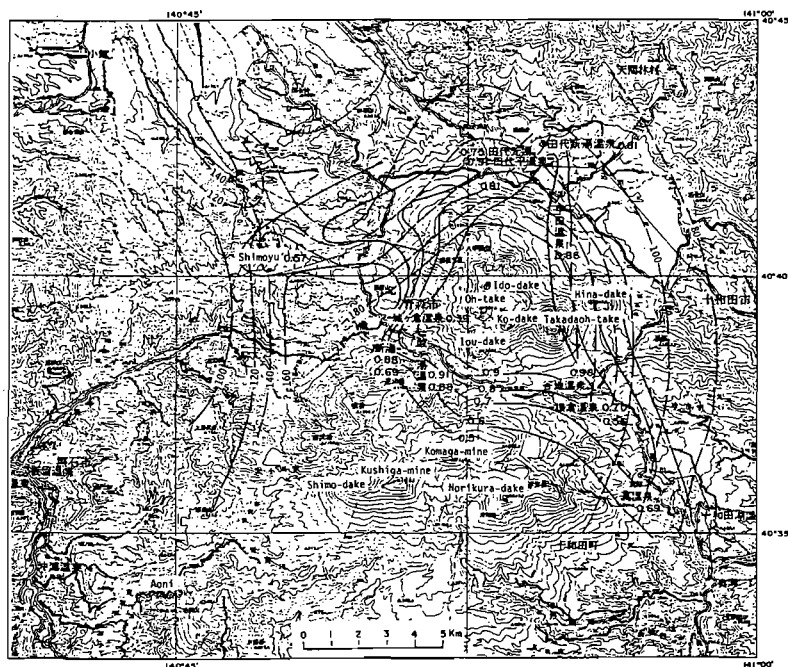
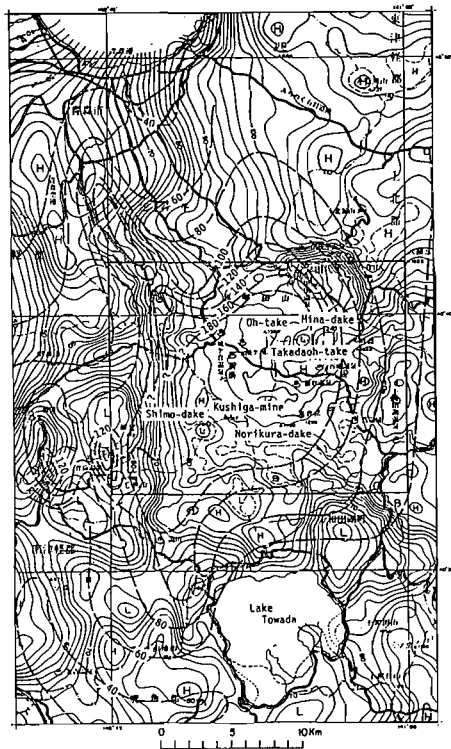


Fig.2 Comparison of anion index map(solid curves; New Energy Development Organization, 1986a) with temperature gradient map(broken curves).



REFERENCES

New Energy Development Organization(1986a):
[Nationwide Survey of Geothermal
Exploration, Report of Geochemical Survey
(Hakkoda area)], 1-59pp.(in Japanese).

New Energy Development Organization(1986b):
[Nationwide Survey of Geothermal
Exploration, Report of Gravity Survey
(Hakkoda area)], 1-113pp.(in Japanese).

Okubo,Y., Tsu,H., and Ogawa,K.(in press):
Estimation of Curie Point Temperature and
Geothermal Structure of Island Arcs of
Japan, Tectonophysics.

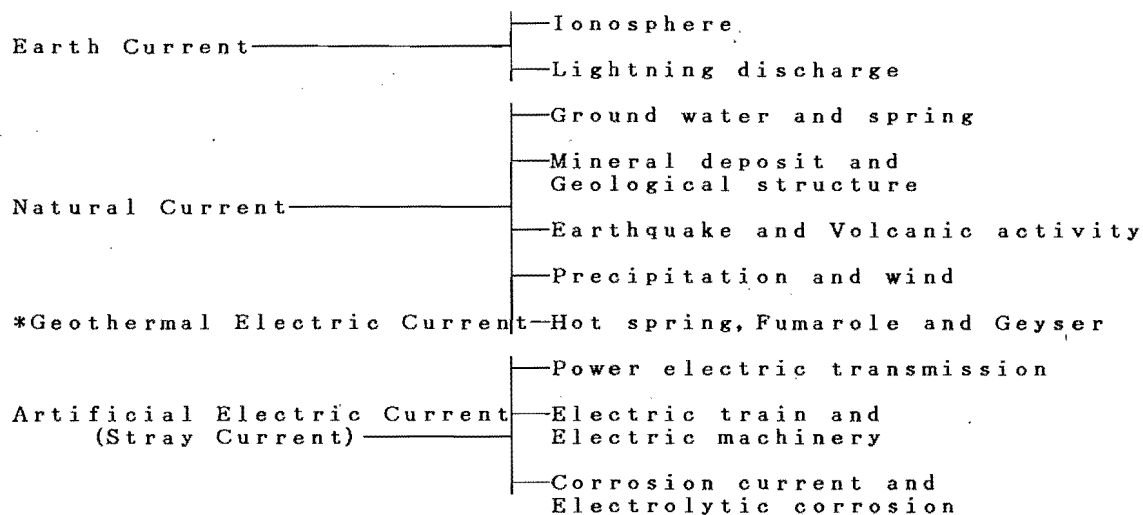
Fig.3 Comparison between Bouguer anomaly reduced at a density of 2.30 g/cm³(solid curves; New Energy Development Organization, 1986b), and temperature gradient(broken curves) Contour values of Bouguer anomaly are in unit of mGal.

GROUND ELECTRIC CURRENT AT GEOTHERMAL AREA

ITO, Y., SHIBATA, A., KIRIHARA, H., NAGUMO, M., and TANAKA, Y.,
Faculty of Eng., KANTO GAKUIN Univ., Yokohama 236, Japan

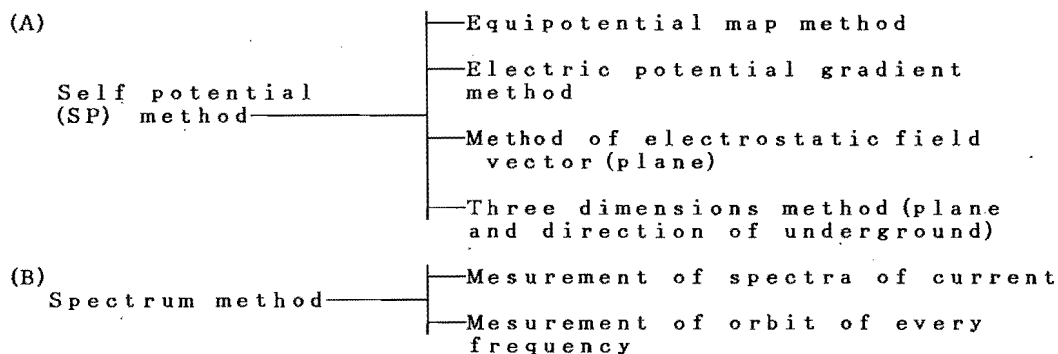
1. The classification of ground electric current

The main factors of generation of the ground electric current are classified as next table.



2 The methods of measurement

The methods of measurement are the next two kinds.



3 Goshogake Geothermal Area, Akita Prefecture, Northern Japan

In Goshogake Geothermal Area, there are a number of boiling springs nonerupting spring, fumaroles and mud pots. And there is a geyser in a hot pool. Estimation of heat discharges and analyse of chemical composition at this area were made by Dr. K. Yuhara, Dr. S. Ehara et al in 1983. Fig. 1 shows the distribution of 1m dpth temperature. And Measurements of geothermal seismic noise wer made in the same area by Dr. S. EHARA et al in 1983.

We made mesurements of geothermal electric current in same geothermal area in 1987 and 1988.

4 Geothermal Electric Current

The causes of generation of geothermal electric current are illustrated with four patterns of Fig. 2.

Theoretically, they are caused by streaming potential, that is, electrostatic field takes rise in stationary flow state of water, vapour and gas, besides, the time variation of geothermal electric current takes rise in geysering action or pulsation of springs and fumaroles.

5 Results of observation

Fig. 3 shows observation points and Fig. 4 shows vectors of electric potential gradient by SP method in geothermal area. Fig. 5 and Fig. 6 show points appeared predominant frequency of spectra of geothermal electric current by spectrum method and orbit of same current, respectively, in geothermal area.

The directions of vector seem to indicate the opposite direction to the high temperature spot in geothermal area. As to spectrum of every points, the predominant frequencies BA, A, B and C type arrange in the order of higher temperature spot. The distribution of orbit of geothermal current is not clear about the physical meaning.

6 Conclusion

The generations of geothermal electric current appear clearly in the geothermal area. The explanation of generation mechanism will bring about a large profit to the prospecting of geothermal area.

References

- Y. ITO, T. SAITO and M. NAGUMO (1978): On the Measurements of the Streaming Potential caused by the Zenikawa Geyser, Part 1, J. G. E. A., Vol. 15, 19-27 (in Japanese)
- Y. ITO, T. SAITO, M. SEKIOKA and K. YUHARA (1983): Telluric currents in the Kurobe Hot Tunnel, Kurobe-Sennindani Hot Dry Rock, J. G. R. S. J., Vol. 5, 39-54 (in Japanese)
- Y. ITO, T. SAITO, K. ICHIKAWA, M. NAGUMO, H. KAWAGUCHI and A. TAKEUCHI (1984): Proposal of a New Technique to Detect the Flow Rate of Underground Water from the Variations of Streaming Potential-The Streaming Potential Technique -, J. J. A. G. W. H., Vol. 26, 77-96 (in Japanese)
- K. YUHARA, S. EHARA, S. AKIBAYASHI and T. NODA (1985): Water and Heat Budgets and Chemical Components of a Hot Pool, Oyunuma, Goshogake Geothermal Area, Northern Japan, J. G. R. S. J., Vol. 7, 131-158 (in Japanese)
- S. EHARA and H. KITAMURA (1985): Geothermal Seismic Noise Measurements in the Goshogake Area, Northern Japan, J. G. R. S. J., Vol. 7, 359-381 (in Japanese)
- Y. ITO, T. SAITO, M. NAGUMO and K. ICHIKAWA (1986): Thermal Properties of Rock and Soil, B. I. T., K. G. Univ., No. 11, 35-49 (in Japanese)
- Y. ITO, T. SAITO, H. KIRIHARA, M. INAGAKI and Y. GOTO (1986): Nanatugama Geyser of Kamisuwa Spa in Suwa City, Nagano Prefecture, Japan, Part I., Bimodality of interval between eruptions and pattern of energy release, J. B. S. J., Vol. 36, 125-136 (in Japanese)

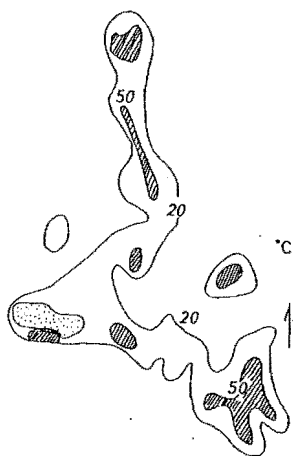


Fig. 1 Distribution of temperatures at 1 m depth observed in Goshogake geothermal area in 1981. Unit: °C

K. YUHARA et al (1985)

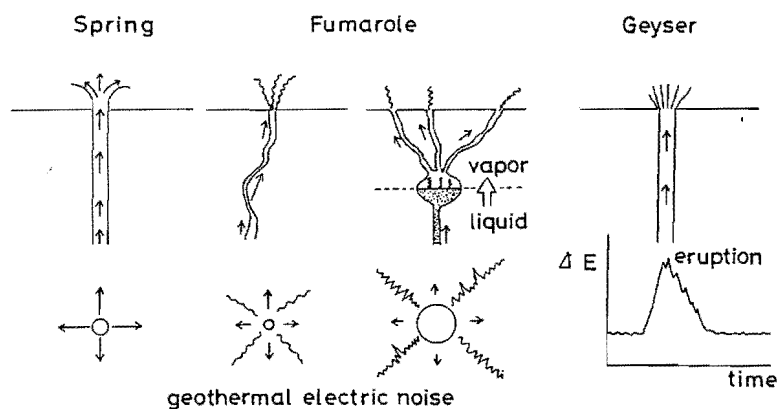


Fig. 2 Generation of, geothermal electric noise

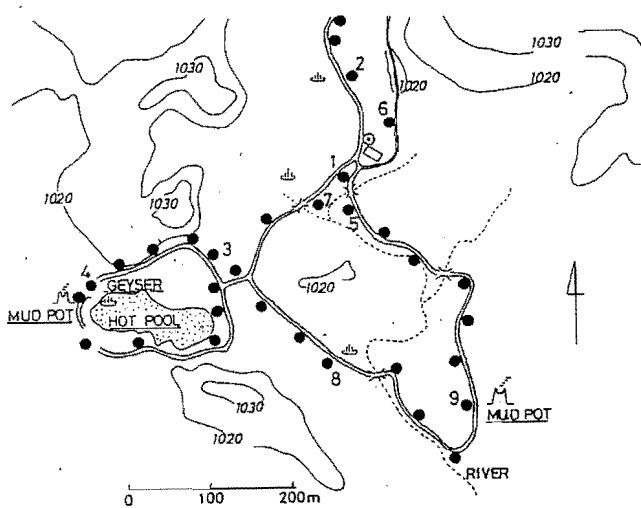


Fig. 3 Observation points in Goshogake geothermal area in 1987 and 1988

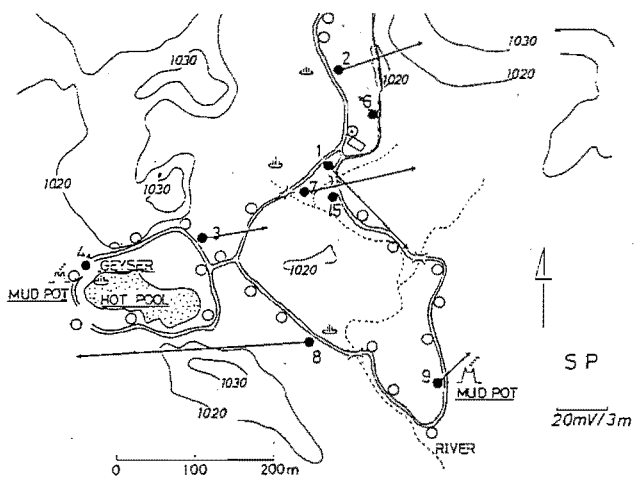


Fig. 4 Vector of SP in 1987

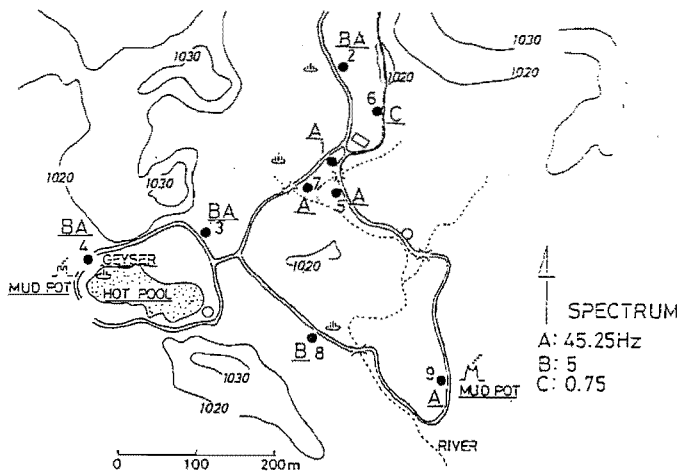


Fig. 5 Geothermal electric spectrum in 1987

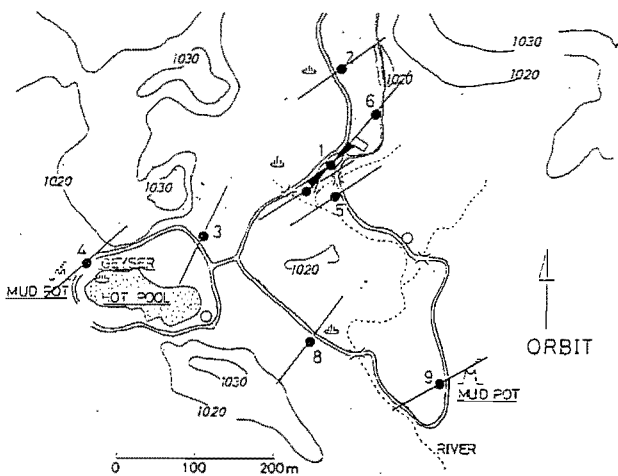


Fig. 6 Orbit of predominant frequency of geothermal electric noise in 1987