







# Fault permeability models for geothermal doublet designs

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#### ABSTRACT

The occurrence and properties of natural faults and fractures in geothermal reservoirs are key in determining reservoir flow properties, and thereby the performance of geothermal doublets placed in fractured reservoirs or in the vicinity of fault zones. In this paper, an analytical model is presented that describes the 3D non-isotropic permeability of a geothermal reservoir around a major fault zone, taking into account typical fault architectures consisting of a fault core, a damage zone and surrounding intact reservoir matrix. The sensitivity of model predictions to orientations of sedimentary layers, damage zone fractures and fault core is analysed for typical permeability contrasts and dimensions of intact reservoir, damage zone fractures and fault core. The model can be used to determine optimum orientation of geothermal doublets around fault zones, taking into account the distribution and characteristics of faults, fractures and sedimentary layering. Implications for optimizing the design of geothermal doublets placed in the vicinity of large fault zones are given.

### 1. INTRODUCTION

The occurrence and properties of natural faults in geothermal reservoirs are key in determining reservoir flow properties, and thereby the success of many geothermal projects (Hickman et al 1997; Fairly and Hinds 2004; Faulds et al. 2010; Moeck 2014). Accordingly, exploration for new geothermal sites will benefit from site-specific data on fault-related factors like damage zone fracture density, connectivity and permeability. In most cases, such data is lacking during geothermal exploration, but existing knowledge can be used to constrain typical fault zone architectures, spatial distribution of permeability and characteristics of damage zone fracture populations.

The characteristics of natural fault zones can vary widely depending on local geological and tectonic settings (Kim et al 2004; Wibberley et al 2008; Faulkner et al 2010). Site-specific characteristics of fault and fracture populations have been determined using seismic surveys, outcrop analogues, core material, and laboratory experiments (Odling et al

1999; Bonnet et al 2001; Torabi and Berg 2011). These studies have provided some generic fault scaling relations that can be used to constrain fault zone characteristics, such as damage zone width, fracture density and dominant fracture orientation. Accordingly, typical fault zone permeability models can be constructed even if site-specific (subsurface) data on fault zone characteristics is lacking.

The relationship between fluid flow and faulting has been studied extensively (Frank 1965; Sibson 1981; Caine et al 1996; Fisher and Knipe 2001). Most fault zones consists of a specific architecture with different structural units, i.e. a single or multiple fault cores and damage zones, surrounded by intact reservoir rock (Chester and Logan 1986; Wibberley et al 2008; Faulkner et al 2010). Permeability may vary considerably in each of these structural units, and the characteristic architecture of a fault zone will determine the permeability in and around fault zones (Barton et al 1995; Wibberley and Shimamoto 2003; Mitchell and Faulkner 2012).

Although many studies that model the permeability around fault zones address some aspects of fault architecture (Yielding et al 1997; Brown and Bruhn 1998; Odling et al 2004), few consider 3D permeability in all structural units. Moreover, one dominant fracture set with equally-spaced fractures is often assumed to describe fractured media such as damage zones (Pickup et al 1995; Lei et al 2015), hence ignoring relations between the orientation of main faults and damage zone fractures (Tchalenko 1970; Kim et al 2004) and spatial variation in fracture density (Mitchell and Faulkner 2012). Another major application to fault-controlled challenge for geothermal energy resources is the general lack of site-specific data on fault and fracture populations.

In this study, an analytical model has been derived for the permeability of a geothermal reservoir around a major fault zone. The model is based on a fractured shale permeability model by Ter Heege (2016), and is modified to be applicable to geothermal energy extraction around fault zones. Fault zones are modelled with a typical architecture consisting of 3 structural units, i.e. a fault core, a damage zone and surrounding intact reservoir matrix. Permeability is modelled using 3D permeability tensors that describe

non-isotropic permeability in each of the structural units. Permeability of intact reservoir rock includes anisotropy due to sedimentary layering. Damage zone permeability is based on fracture density and orientation, and includes the effect of decreasing fracture density with increasing distance from the fault core. The permeability of the fault core can include anisotropy due to fabric in the fault gouge. Bulk permeability can be described by combining permeability of the fault core, damage zone and intact reservoir including typical dimensions of each structural unit as well as distance and orientation between geothermal doublets.

The bulk permeability model can be used to analyse the geothermal power for a doublet system consisting of a surface heat exchanger, an injection well ("injector") and a production well ("producer") that is placed in the vicinity of a fault zone. Implications for the optimum design of geothermal doublet systems placed around fault zones are discussed. It is shown that produced geothermal power can be considerably enhanced if doublets are placed in optimum orientation with respect to existing damage zone fracture populations. Therefore, optimization of geothermal doublet designs in geothermal reservoirs that are crosscut by fault zones help de-risking geothermal exploration and exploitation.

# 2. MODEL FOR FAULT-CONTROLLED RESERVOIR PERMEABILITY

Fault zones are modelled with a typical architecture consisting of 3 structural units, i.e. a fault core, a damage zone and surrounding intact reservoir matrix (Fig. 1, c.f. Ter Heege 2016). Dimensions of fault zone and its structural components are constrained by length along the fault strike (L), length along fault dip (S), displacement along fault dip  $(D_S)$ , damage zone width  $(W_D)$ , damage zone fracture width  $(W_F)$ , and fault core width  $(W_C)$ . Anisotropic permeability in each of the units is described using 3D permeability tensors (K) that are defined in terms of 3 orthogonal principal permeabilities in different coordinate systems.

$$\boldsymbol{K}^* = \begin{pmatrix} K_{11}^* & 0 & 0 \\ 0 & K_{22}^* & 0 \\ 0 & 0 & K_{33}^* \end{pmatrix}$$
[1]

The asterisk in equation [1] is included to acknowledge that the principal permeabilities are defined using different Cartesian coordinate systems for the different structural units (following the right-hand rule for axes orientations).

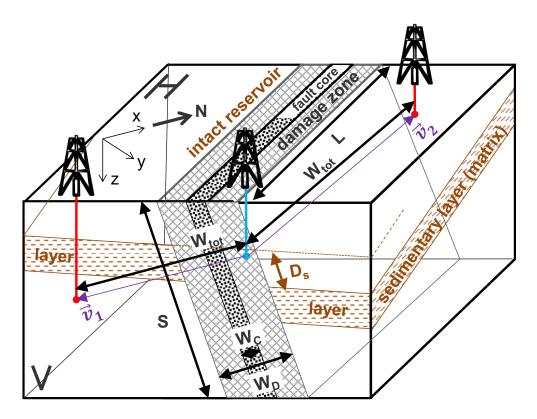


Figure 1: Schematic diagram of simplified fault zone showing the structural units (i.e. non-fractured intact shale matrix with sedimentary layers, damage zone and fault core), and dimensions used in the fractured shale permeability model (L- length along fault strike, S- length along fault dip,  $D_{S}$ - displacement along fault dip,  $W_{D}$ - damage zone width,  $W_{C}$ - fault core width,  $W_{tot}$  – distance between injector and producer doublet wells,  $v_{x}$ - orientation of geothermal doublet injector-producer). The geographical coordinate system (axis x, y, z in northern, eastern and depth direction, respectively) is also indicated.

Permeability tensors are transformed between the different coordinate systems by tensor transformations using a rotation matrix R for a full Euler rotation based on the orientations of principal permeabilities relative to matrix layers, fault core and damage zone.

The matrix permeability tensor  $(\mathbf{K}^* = \mathbf{K}_{\mathbf{M}}')$  is defined on the basis of sedimentary layers that may be present in the intact reservoir at an orientation described by the dip direction  $(\theta_M)$  and dip  $(\varphi_M)$ . The following orientations of principle permeabilities are assumed: (1)  $K_{MII}$ ' is within the layers at an angle  $\omega_M$  relative to  $\theta_{M}$ , (2)  $K_{M22}$ ' is within the layers perpendicular to  $K_{M11}$ ', and (3)  $K_{M33}$ ' is perpendicular to the layers in downward direction. Together, the permeabilities define a layer-based coordinate system. The matrix permeability tensor and layer-based coordinate system can account for effects of sedimentary structures such as cross-bedding (i.e. by specifying  $\omega_M$  and  $K_{M11}'/K_{M22}'$ ) and permeability anisotropy due to bedding (i.e. by specifying  $K_{M11}$  '/ $K_{M33}$  ').

The damage zone permeability tensor  $(\mathbf{K}^* = \mathbf{K}_D)$  is defined on the basis of a dominant fracture set that may be present in the damage zone at an orientation described by the dip direction  $(\theta_F)$  and dip  $(\varphi_F)$ . The following orientations of principle permeabilities are assumed: (1)  $K_{FII}$  "is parallel to the normal vector of fracture planes for the dominant fracture set, (2)  $K_{F22}$ " is parallel to the layer strike, and (3)  $K_{F33}$ " is parallel to the layer dip vector. Together, the principle permeabilities define a fracture-based coordinate system. The principle damage zone permeabilities are expressed as a combination of matrix permeability  $(K_M)^{**}$ , now in the fracture-based coordinate system) and damage zone fracture permeability  $(K_F)^{**}$ . The damage zone permeability tensor and fracture-based coordinate system can account for effects of permeability anisotropy in the fractures (i.e. by specifying  $K_{F11}$ ''',  $K_{M22}$ ''' and  $K_{M33}$ '''), different fracture orientations (i.e. by specifying  $\theta_F$  and  $\varphi_F$ ), and different fracture densities. Damage zone permeability perpendicular to the fracture plane  $(K_{DII})$  is calculated using the harmonic mean of diagonal tensor components ( $K_{MII}$ ''' and  $K_{FII}$ '''), while damage zone permeability parallel to the fracture plane ( $K_{D22}$ ) and  $K_{D33}$ ''') is calculated by the arithmetic mean of  $K_{M22}$ ''' and  $K_{F22}$ ''' and  $K_{M33}$ ''' and  $K_{F33}$ ''', respectively. As off-diagonal tensor components of  $K_F$ ", are zero per definition, the off-diagonal components of  $K_D$ " that govern cross-flow are based on  $K_M$ " (e.g., Pickup et

$$K_{D11}^{"} = \left(\frac{F_{w}W_{F}}{K_{F11}^{"}} + \frac{\left(1 - F_{w}W_{F}\right)}{K_{M11}^{"}}\right)^{-1}$$
 [2a]

$$K_{D22}^{""} = F_w W_F K_{F22}^{""} + (1 - F_w W_F) K_{M22}^{""}$$
 [2b]

$$K_{D33}^{""} = F_w W_F K_{F33}^{""} + (1 - F_w W_F) K_{M33}^{""}$$
 [2c]

$$K_{Dij}^{"'} = (1 - F_w W_F) K_{Mij}^{"'} \text{ for } i \neq j$$
 [2d]

Where the contribution of matrix and damage zone fractures to damage zone permeability is weighted using damage zone fracture width  $(W_F)$  and density of fractures in the damage zone  $(F_w)$ . Fracture density decreases with increasing distance (w) from the center of the fault core (c.f. Mitchell and Faulkner, 2012).

$$F_{w} = F_{C} \exp\left(\frac{w - W_{C}}{w - W_{D}}\right) \text{ with } W_{c} \le w \le W_{D} \quad [3]$$

Where distance from the fault core is normalized to give fracture density  $F_w = F_C$  at the outer edge of the fault core  $(w = W_C)$  where permeability equals that of the fault gouge, and  $F_w \to 0$  at the outer edge of the damage zone  $(w \to W_D)$  where permeability equals that of the matrix. It is assumed that one dominant set controls permeability, but the model may be extended to include multiple fracture sets as frequently observed in fault zones (Tchalenko, 1970; Kim et al., 2004).

The fault core permeability tensor ( $\mathbf{K}^* = \mathbf{K}_C$ ") is defined on the basis of the dominant fault zone orientation described by the dip direction ( $\theta_C$ ) and dip ( $\phi_C$ ) of the fault core. The following orientations of principle permeabilities are assumed: (1)  $K_{C11}$ " is parallel to the normal of the main fault plane through the fault core, (2)  $K_{C22}$ " is parallel to the strike of the main fault plane, and (3)  $K_{C33}$ " is parallel to the dip vector of the main fault plane. Together, the principle permeabilities define a fault-based coordinate system. The fault core permeability tensor and fault-based coordinate system can account for effects of permeability anisotropy in the fault core (i.e. by specifying  $K_{C11}$ ",  $K_{C22}$ " and  $K_{C33}$ ").

The principle bulk permeabilities around a single fault zone are determined in the fault-based coordinate system ( $K_B$ ') by combining matrix permeability ( $K_M$ '), average damage zone permeability ( $K_D$ ') and fault core permeability ( $K_C$ '), averaged over the distance ( $W_{tot}$ ) between the doublet injector and producer. Equations [2] give the permeability tensor components for the damage zone at a distance from the center of the fault core.  $K_D$ ' is calculated based on the average permeability over distance  $w_D = w_{D2} - w_{DI}$  in the damage zone that can be calculated by substituting equation [3] in [2] and integrating local permeability tensor components over  $w_D$ .

$$\tilde{K}_{Dij}^{"''} = \frac{1}{w_D} \int_{w_{D1}}^{w_{D2}} K_{Dij}^{"'} dw$$
 [4]

Where  $w_D$  is determined by the location of the injector and producer relative to fault core and damage zone (c.f. Fig. 1). Bulk permeability perpendicular to the fault core  $(K_{BII})$  is calculated using the harmonic mean of diagonal tensor components  $K_{MII}$ ,  $K_{DII}$  and  $K_{C33}$ . Bulk permeability parallel to the fracture plane  $(K_{B22})$  and  $K_{B33}$  is calculated by the arithmetic mean of  $K_{M22}$ ,  $K_{D22}$ ,  $K_{C22}$ , and  $K_{M33}$ ,

 $K_{D33}$ ",  $K_{C33}$ ", respectively. As off-diagonal tensor components of  $K_C$ " are zero per definition, the off-diagonal components of  $K_B$ " that govern cross-flow are based on  $K_M$ " and  $K_D$ ".

$$K_{B11}^{"} = \left(\frac{\left(W_{tot} - W_D - W_C\right)}{W_{tot}K_{M11}^{"}} + \frac{W_D}{W_{tot}K_{D11}^{"}} + \frac{W_C}{W_{tot}K_{C11}^{"}}\right)^{-1}$$

[5a]

$$K_{B22}^{"} = \frac{(W_{tot} - W_D - W_C)K_{M22}^{"}}{W_{tot}} + \frac{W_D K_{D22}^{"}}{W_{tot}} + \frac{W_C K_{C22}^{"}}{W_{tot}}$$
[5b]

$$K_{B33}^{"} = \frac{\left(W_{tot} - W_D - W_C\right)K_{M33}^{"}}{W_{tot}} + \frac{W_DK_{D33}^{"}}{W_{tot}} + \frac{W_CK_{C33}^{"}}{W_{tot}}$$
[5c]

$$K_{B1j}^{"} = \left(\frac{\left(W_{tot} - W_D - W_C\right)}{W_{tot}K_{M1j}^{"}} + \frac{W_D}{W_{tot}K_{D1j}^{"}}\right)^{-1} \text{ for } j = 2, 3$$
[5d]

$$K_{Bij}^{"''} = \frac{(W_{tot} - W_D - W_C)K_{Mij}^{"'}}{W_{tot}} + \frac{W_D K_{Dij}^{"}}{W_{tot}} \quad i \neq j; \quad i = 2, 3$$
[5e]

The tensor for average bulk permeability  $(K = K_B)$  is defined with the following orientations of principle permeabilities: (1)  $K_{B11}$  is South to North, (2)  $K_{B22}$  is West to East, and (3)  $K_{B33}$  is downward along depth. Together, the principle permeabilities define a geographical based coordinate system. Dip directions  $(\theta_M, \theta_C, \theta_F)$  and dip angles  $(\varphi_M, \varphi_C, \varphi_F)$  are defined within the geographical coordinate system.  $K_B$  can be derived from  $K_B$ " by transformation between the faultbased and geographic coordinate system using a rotation matrix that is based on fault dip direction ( $\theta_c$ ) and dip angle  $(\varphi_C)$ . Bulk permeability in the direction between the injector and producer of a doublet (in the geographical coordinate system) can be described by a permeability vector with three components parallel to S-N, W-E and along depth (x-, y- and z-directions, respectively). This vector can be calculated by multiplying  $K_R$  with the unit vector ( $\vec{v}$ ) describing the direction between the injector and producer. The average bulk permeability value that can be used to calculate geothermal power for a doublet system is given by the magnitude of the permeability vector, for example following the method outlined by Van Wees et al (2012) for isotropic permeability. It should be emphasized that the current model incorporates fault zone architecture in the description of anisotropic permeability of geothermal reservoirs. Thereby, it can be used to search for optimum doublet designs that yield the best performance in terms of geothermal energy extraction.

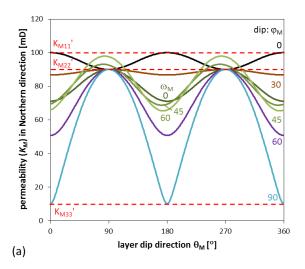
# 3. MODEL SENSITIVITY TO ORIENTATIONS OF LAYERS, FRACTURES AND FAULTS

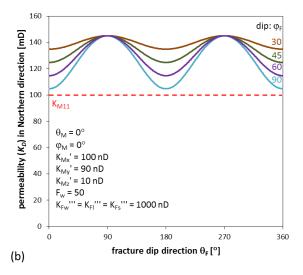
Model predictions are given for typical permeability contrasts and dimensions of intact reservoir, damage zone fractures and fault core (Table 1).

Table 1: Input parameters for model sensitivity analysis

Parameter	Unit	Value
K <sub>M11</sub> '	mD	100
K <sub>M22</sub> '	mD	90
K <sub>M33</sub> '	mD	10
K <sub>F11</sub> '''	mD	1000
K <sub>F22</sub> '''	mD	1000
K <sub>F33</sub> '''	mD	1000
K <sub>C11</sub> ''	mD	2000
K <sub>C22</sub> ''	mD	2000
K <sub>C33</sub> ''	mD	2000
$W_{\mathrm{F}}$	m	0.001
$W_{\mathrm{D}}$	m	100
$W_{\rm C}$	m	0.1
W <sub>tot</sub>	m	1000
$F_{C}$	#/m	100

Orientations of fault core, damage zone fractures and reservoir bedding or layering are varied to show the sensitivity of permeability to fault and fracture orientations for each of the structural units (Fig. 2). The example shows values for bulk permeability, assuming that the doublet is oriented in S-N direction ( $\vec{v} = [1,0,0]$ ), and placed in the intact reservoir matrix outside the damage zone or in the damage zone. Accordingly, the variation in bulk permeability results from permeability anisotropy due to layering and density of damage zone fractures, and different orientations of the layers and fractures relative to the doublet.





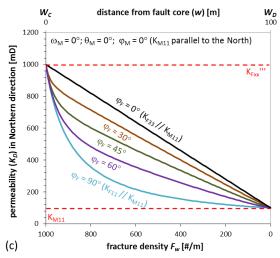


Figure 2: Variation of bulk permeability for a doublet oriented in S-N direction for: (a) different orientations of sedimentary layering in the reservoir (given by layer dip direction  $\theta_M$  and dip angle  $\varphi_M$ ) and largest principal matrix permeability within the matrix layers ( $\omega_M$ ), (b) different orientations of damage zone fractures (given by fracture dip direction  $\theta_D$  and dip angle  $\varphi_D$ ) and  $F_w =$ 50, (c) increasing distance from fault core (w) and associated decrease in fracture density  $(F_w, \text{ c.f. Eq. 6})$ . Model input parameters for the sensitivity analysis are given in Table 1.

# 4. IMPLICATIONS FOR OPTIMIZATION OF GEOTHERMAL DOUBLET DESIGNS AROUND FAULT ZONES

Fluid flow around a fault zone crosscutting a geothermal reservoirs is typically controlled by damage zone fractures or deformation bands (Fossen et al 2007; Mitchell and Faulkner 2012). These structures may give rise to high permeability pathways, for example in geothermal reservoirs with low matrix permeability (e.g., tight sandstone or

rocks) when local deformation is crystalline accompanied by dilatation. The permeability contrast used in the example used to analyze model sensitivity (c.f. Fig. 2) is typical of this situation. In the example, high permeability pathways are oriented parallel to the fracture planes. The sensitivity analysis shows that for a doublet oriented in S-N direction, permeability is highest if fractures dip to the West or East (Fig. 2b). The maximum difference in permeability is a factor 1.5 for fractures dipping 90° to the West or East compared to fractures dipping 90° to the North or South in the example. In general, permeability is higher for shallow dipping fractures compared to steeper dipping fractures. Maximum permeabilities occur close to the fault core (Fig. 2c). In contrast, the opposite trends may occur in a geothermal reservoir with high matrix permeability (e.g., porous sandstone or carbonate rocks) where local deformation may be accompanied by compaction resulting in bands of low permeability. If the doublet is placed outside of the damage zone (at  $w > 0.5W_D$  from the fault core), permeability difference is mainly determined by the anisotropy caused by the sedimentary layering (Fig.

If dominant damage zone fractures are at Riedel orientation (R-shear), high permeability pathways resulting from the fractures are oriented at an angle of  $45^{\circ}$  -  $\varphi/2$  relative to the fault core (assuming it aligns with direction of maximum shear displacement), and at an angle of - $\varphi/2$  relative to the maximum principal stress (c.f. Ter Heege 2016). Hence, structural interpretation of fault patterns or analysis of the stress field aid in determining the optimum orientation of geothermal doublets.

The distribution and characteristics of faults, fractures and sedimentary layering are controlling the performance of geothermal doublets. The optimum orientation of geothermal doublets is dependent on the orientation of sedimentary layers, damage zone fractures and fault core. The presented model, therefore can help optimizing the design of geothermal doublets placed in the vicinity of large fault zones.

It should be emphasized that the choice of optimum drilling location and direction is generally not only motivated by achieving optimum doublet performance, but also by reducing risks of associated with the presence of large faults and fracture networks. Risks may include loss of wellbore stability, screen outs and induced seismicity. Accordingly, optimum orientation of geothermal doublets need to consider optimum doublet performance as well as risks of operations.

## 3. CONCLUSIONS

In this paper, the permeability of geothermal reservoir around fault zones is studied. The main conclusions can be summarized as follows:

 An analytical model of fault-controlled reservoir permeability is derived that describes

- the permeability using 3D permeability tensors for intact reservoir, damage zone and fault core.
- The sensitivity of the model to the orientations of layers, fractures and faults as well as damage zone fracture density is analyzed. It shows that bulk permeability can vary considerably, mainly depending on (i) permeability anisotropy in the matrix, damage zone and fault core, (ii) the orientation of matrix layers and damage zone fractures, and (iii) the location relative to the fault core.
- The model can be used to determine optimum orientation of geothermal doublets around fault zones, taking into account the distribution and characteristics of faults, fractures and sedimentary layering. Thereby, it helps optimizing the design of geothermal doublets placed in the vicinity of large fault zones.

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